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GRAVITY SURVEY - HAMLIN VALLEY

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# Prepared for:

U.S. Department of the Air Force Ballistic Missile Office (BMO) Norton Air Force Base, California 92409

# Prepared by:

Fugro National, Inc. 3777 Long Beach Boulevard Long Beach, California 90807

7 February 1980

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#### **FOREWORD**

Methodology and Characterization Studies during fiscal years 1977 and 1978 included gravity surveys in ten valleys in Arizona (five), Nevada (two), New Mexico (two), and California (one). The gravity data were obtained for the purpose of estimating alluvial gross structure and shape of the basins and the thickness of the valley fill. There was also the possibility of detecting shallow rock in areas between boring locations. Generalized interpretations from these surveys were included in Fugro National's Characterization Reports (FN-TR-26a through e).

During the FY 77 surveys, the measurements were made to form an approximate one-mile grid over the study areas, and contour maps showing interpreted depth to bedrock were made. In FY 79, the decision was made to concentrate the available funds on the basic Verification Program to verify and refine suitable area boundaries. This decision resulted in a reduction in the gravity program. Instead of obtaining gravity data on a grid, the reduced program consisted of obtaining gravity measurements along profiles across the valleys where Verification Studies were also performed.

The Defense Mapping Agency (DMA), St. Louis, was also requested to provide gravity data from their library to supplement the gravity profiles. For Big Smoky, Reveille, and Railroad valleys, a sufficient density of library data is available to permit construction of interpreted contour maps instead of two-dimensional cross sections.

In late summer of FY 79, supplementary funds became available to begin data reduction. At this time, inner zone terrain corrections began on the library data and the profiles from Big Smoky Valley, Nevada, and Butler and La Posa valleys, Arizona. The profile data from Whirlwind, Hamlin, Snake East, White River and Garden Coal valleys, Nevada were available from the field in early October, 1979.

A continuation of gravity interpretations has been incorporated into the FY 80 contract and the results are being summarized in a series of valley reports. The reports covering Nevada-Utah gravity studies will be numbered, "FN-TR-33-", followed by the abbreviation for the subject valley. In addition, more detailed reports of the results of FY 77 surveys in Dry Lake and Ralston valleys, Nevada are being prepared. Verification Studies are continuing in FY 80 and gravity studies are included in the program. DMA will continue to obtain the field measurements and it is planned to return to the grid pattern. The interpretation of the grid data will allow the production of contour maps which will be valuable in the deep basin structural analysis needed for computer modeling in the Water Resources Program. gravity interpretations will also be useful in the Nuclear Hardness and Survivability (NH&S) evaluations.

The basic decisions governing the gravity program are made by BMO following consultation with TRW Inc., Fugro National and the (DMA). Conduct of the gravity studies is a joint effort between DMA and Fugro National. The field work, including planning, logistics, surveying, and meter operation is done by the Defense Mapping Agency Hydrographic/Topographic Center (DMAHTC), headquartered in Cheyenne, Wyoming. DMAHTC reduces the data to Simple Bouguer Anomaly (see Section Al.4, Appendix Al.0). The Defense Mapping Agency Aerospace Center (DMAAC), St. Louis, calculates outer zone terrain corrections.

Fugro National provides DMA with schedules showing the valleys with the highest priorities. Fugro National also recommended locations for the profiles in the FY 79 studies within the constraints that they should follow existing roads or trails. Any required inner zone terrain corrections are calculated by Fugro National prior to making geologic interpretations.

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12 Interpreted Gravity Profile HV-9 13 Interpreted Fault Relationships

#### 1.0 INTRODUCTION

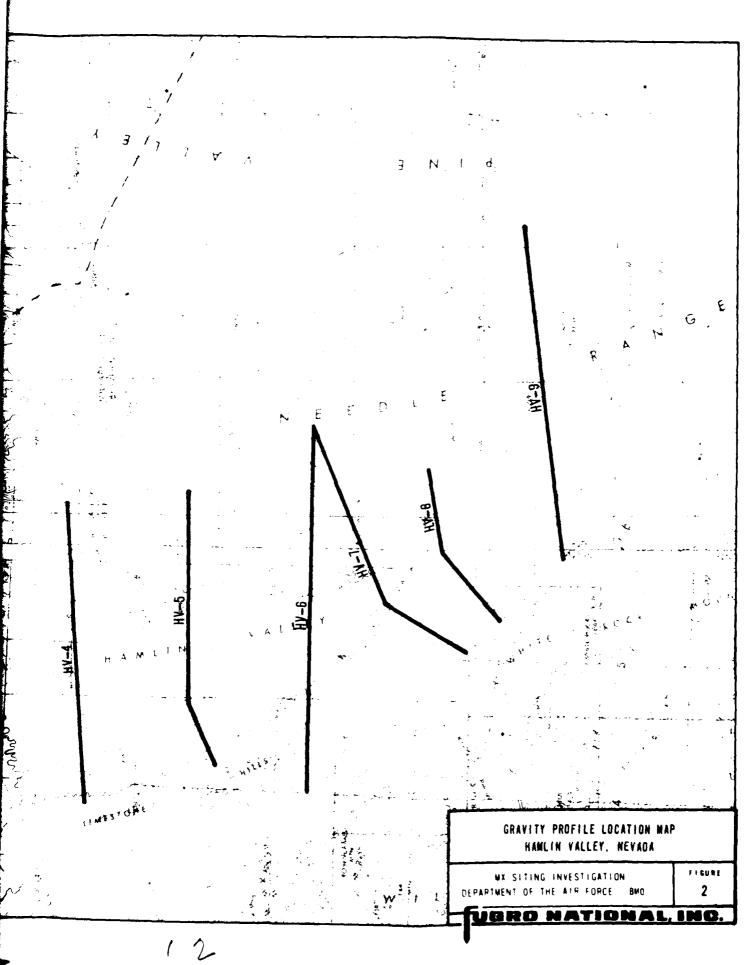
### 1.1 OBJECTIVE

Gravity measurements were made in Hamlin Valley for the purpose of estimating the overall shape of the structural basin and the thickness of alluvial fill. The estimates will be useful in modeling the dynamic response of ground motion in the basin and in evaluating ground-water resources.

#### 1.2 LOCATION

Hamlin Valley is located principally in Lincoln and White Pine counties in east-central Nevada with portions in Millard, Beaver, and Iron counties, Utah (Figure 1). Hamlin Valley is approximately 180 miles (300 km) NNE of Las Vegas, Nevada. U.S. Highway 50 crosses the north end of the valley. There are no paved roads within the valley, but access is generally good along well maintained ranch and mine roads. The nearest towns are Baker, Nevada and Garrison, Utah, along Highway 73 near the northwest corner of the valley.

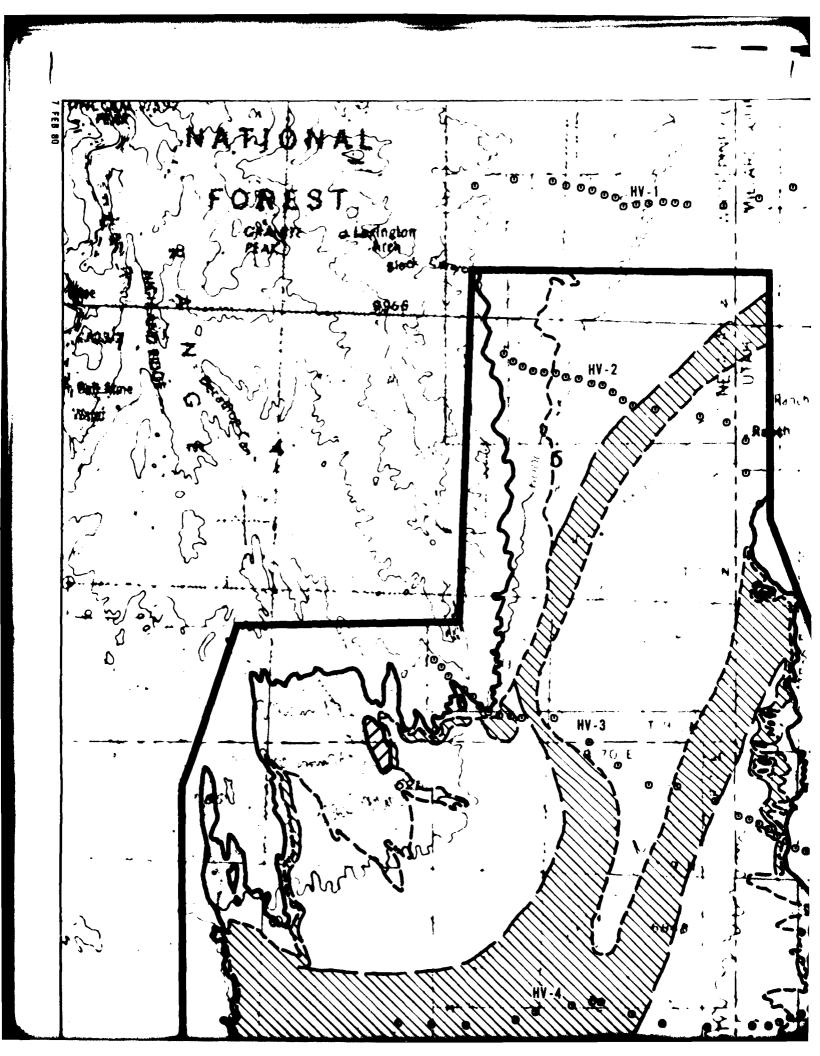
Hamlin Valley is bounded by mountain ranges on three sides and open to Snake Valley on the north (Figure 2). The Burbank Hills and the Needle Range form the eastern boundary of the valley. The Snake Range and the Limestone Hills comprise the western side, and the White Rock Mountains form the southern border. Most of the valley is undeveloped rangeland with several ranches and agricultural fields along Hamlin Wash.

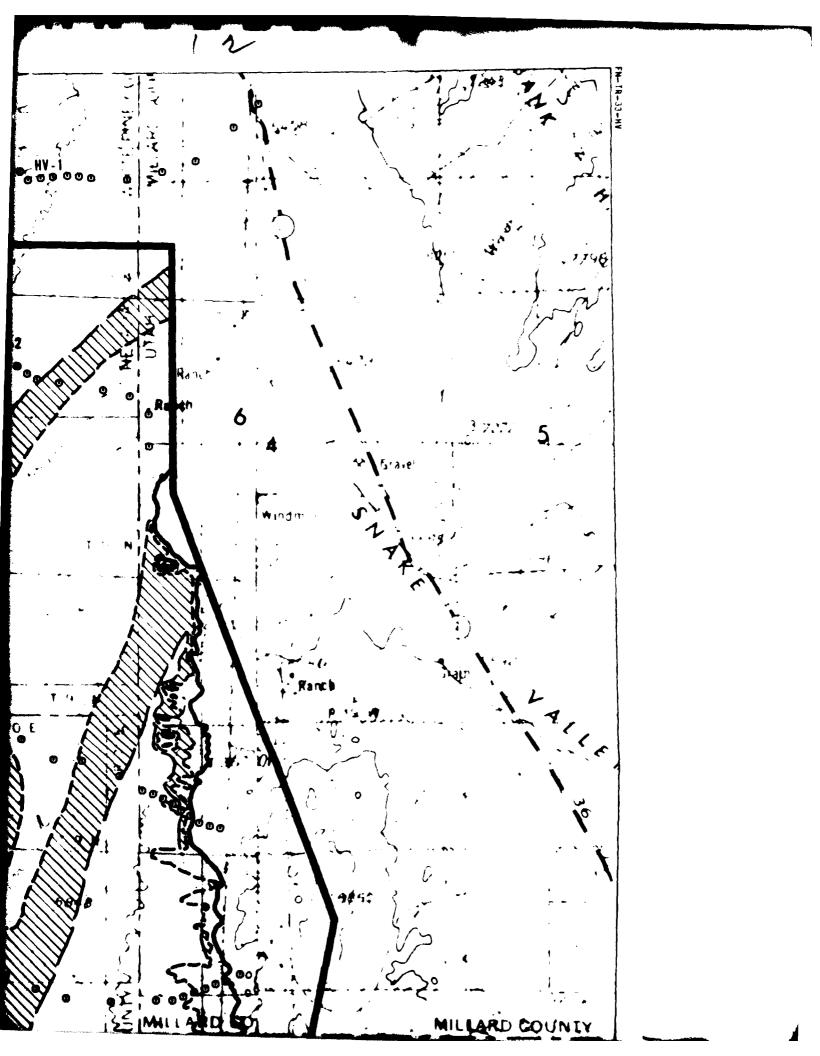


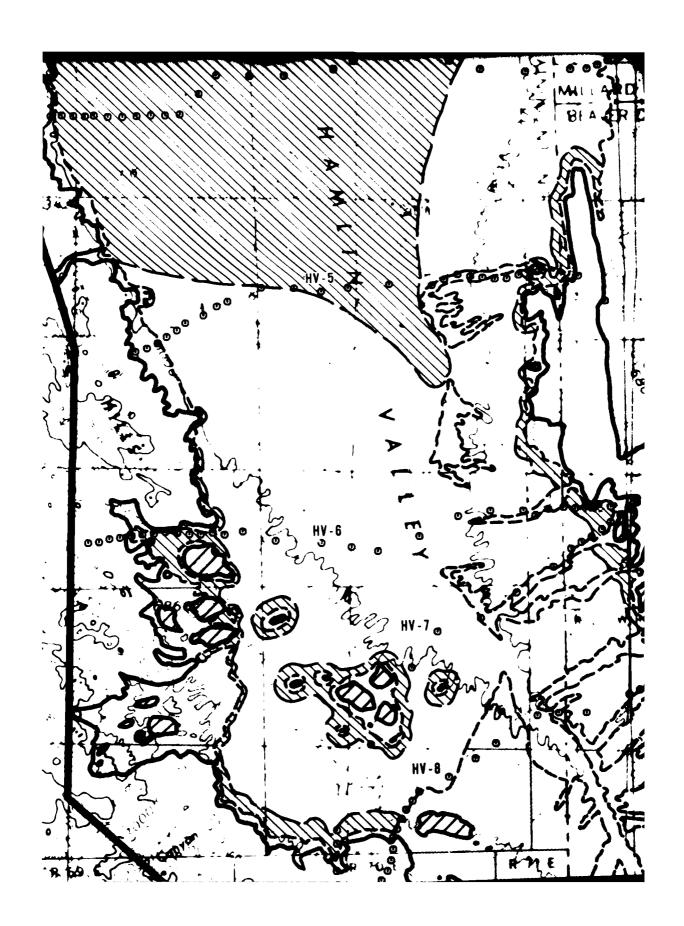
## 1.3 SCOPE OF STUDY

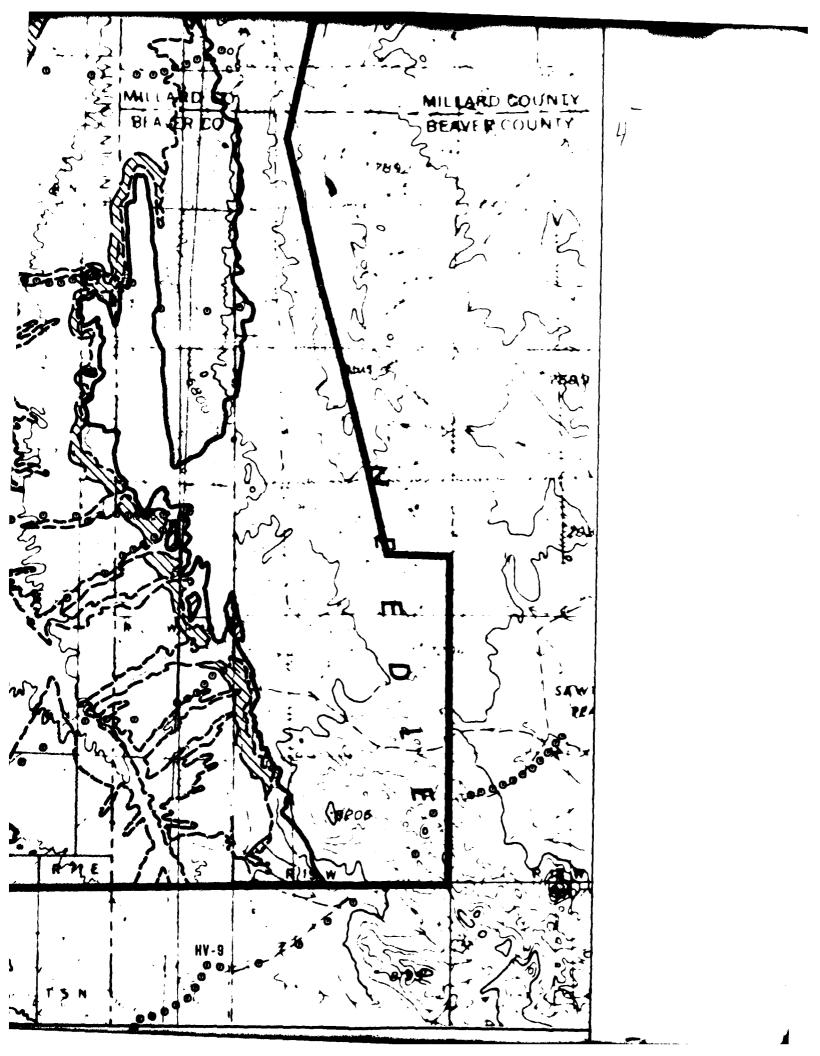
The Defense Mapping Agency Hydrographic/Topographic Center/-Geodetic Survey Squadron (DMAHTC/GSS) obtained 255 gravitational field measurements along nine cross valley profiles in Hamlin Valley. Profile positions are shown in Figure 2 and the locations of the individual stations are shown in Figure 3. The profiles are oriented approximately east-west. They are between 6 and 14 miles (11 to 23 km) long and are spaced approximately 5 miles (8.5 km) apart. The gravity sampling interval was 1 mile (1.6 km) over the central valley section and .25 mile (.4 km) near the valley boundaries. The more dense sampling was used on the valley flanks to define any steep gravity gradients associated with boundary faults and to resolve anomalies with high spatial frequency that could be associated with shallow bedrock.

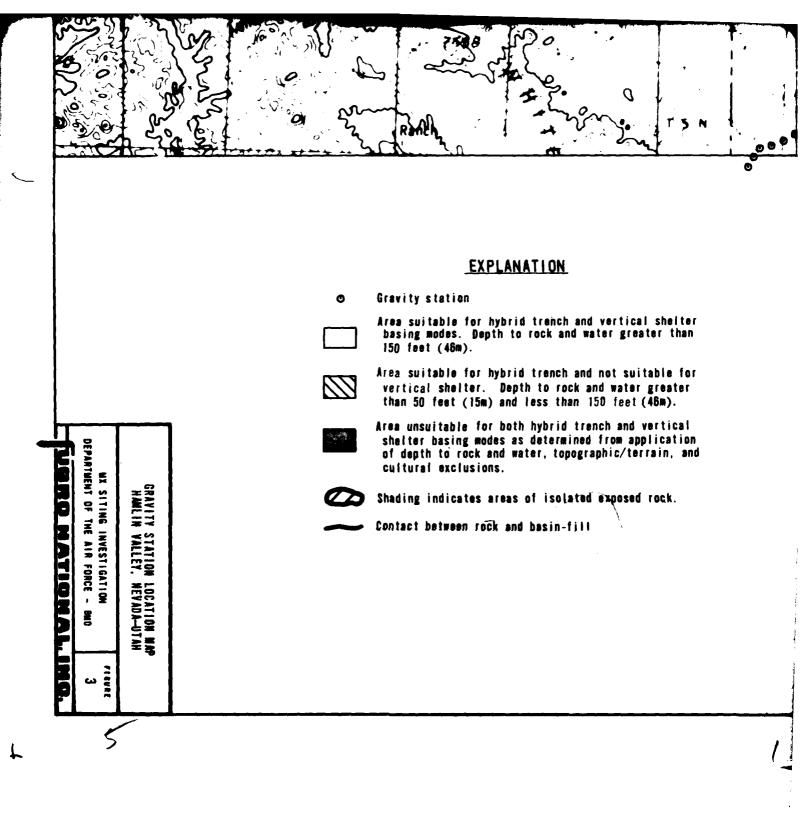
The tolerance for establishing the station elevations was 5 feet (1.5 m), which limits the gravity precision to .3 milligals.

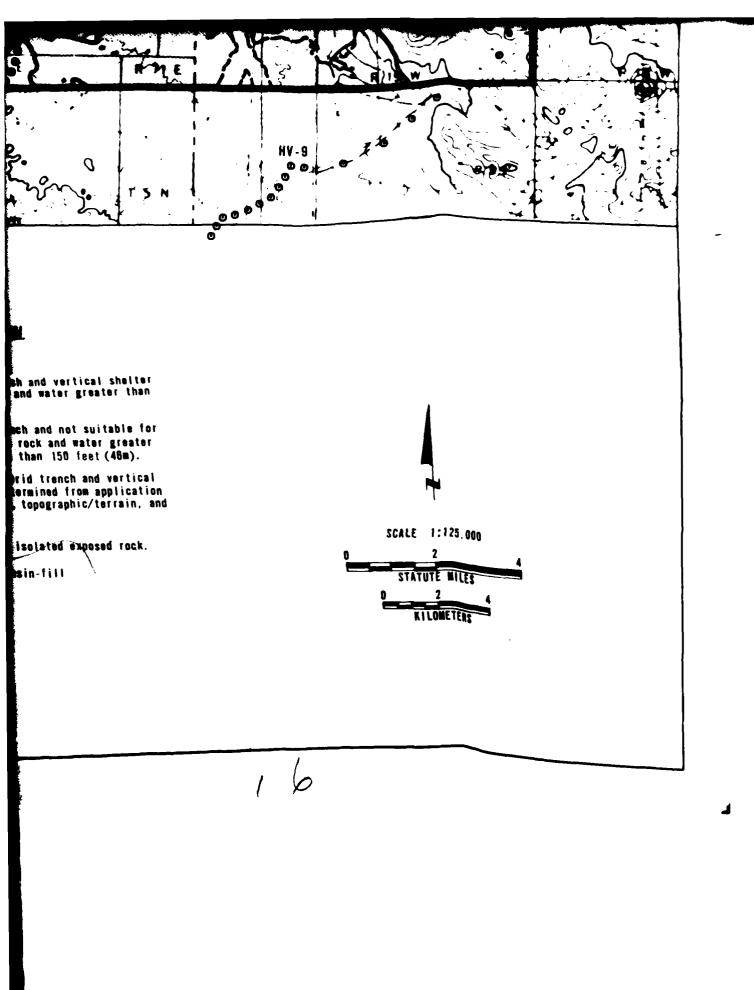












### 2.0 GRAVITY DATA REDUCTION

DMAHTC/GSS obtained the basic observations and reduced them to Simple Bouquer Anomalies (SBA) for each station as described in Appendix Al.O. Up to three levels of terrain corrections were applied to convert the SBA to the Complete Bouguer Anomaly (CBA). First, the Defense Mapping Agency Aerospace Center (DMAAC), St. Louis, used its library of digitized terrain data and a computer program to calculate terrain corrections out to 104 miles (167 km) from each station. When the program could not calculate the terrain effects near a station, a ring template was used to estimate the effect of terrain within approximately 3000 feet of the station. The third level of terrain corrections was applied to those stations where 10 feet or more of relief was observed within 130 feet. In these cases, the elevation differences were measured in the field at a distance of 130 feet along six directions from the stations. These data were used to calculate the effect of the very near relief. CBA data for the Hamlin Valley stations are listed in Appendix A2.0.

# 3.0 GEOLOGIC SUMMARY

The structural geologic setting, major rock types, and depositional regime of the valley fill material are important considerations in the interpretation of gravitational field data.

Hamlin Valley is an elongate north-south trending alluvial basin exhibiting typical Basin and Range structure. Block faulting has formed this strong north-south physiographic framework. According to Hintze (1963), a fault concealed by alluvium parallels the eastern margin of the valley. Rocks in the Needle Range are cut by predominantly north-south trending faults. Most faults in the Snake Range, White Rock Mountains, and Limestone Hills have northwest or north-south orientations.

There is evidence of recent upwarping of the southern end of the valley. The southwestern lake beds are at a relatively high topographic level and are being actively eroded, and the drainage trends anomalously to the north. The alluvium is offset by numerous small faults in this part of the valley (Fugro National, 1979).

The mountains in the northern section of the valley are composed of an assemblage of Paleozoic limestones, dolomites, shales, and sandstones (Hintze, 1963; Hose and Blake, 1976; Tschanz and Pampeyan, 1970). The White Rock Range, on the south and the central portion of the Needle Range, on the east, are composed of undifferentiated Tertiary volcanic rocks.

Most of the surficial deposits in Hamlin Valley are late Tertiary and Quaternary alluvial fan and lake deposits. Basin fill deposits are described in the Verification Studies (FY 79, FN-TR-27-1A). Three relative ages of alluvial fans are recognized. The older fan deposits, consisting of well-indurated gravels and sandy gravels near the Needle and Snake Ranges, are the least extensive unit. Intermediate age alluvial fan deposits are the most extensive surficial unit. They are composed chiefly of silty to clayey sands. Younger alluvial fan, terrace, and modern stream channel deposits consist dominantly of silty sand, and generally occcur in the central part of the valley. Lacustrine deposits, primarily clayey sands, underlie about one-fifth of the valley and are generally restricted to the west side of Hamlin Wash (Fugro National, 1979).

# 4.0 INTERPRETATION

A negative gravity anomaly over a valley such as Hamlin Valley is created when low density alluvial material filling the valley is surrounded by dense rock in the mountains. Gravity profiles across such a valley are often U-shaped, low in the middle where the fill is thickest, and high on the ends where the fill thins and disappears. Interpretation requires removal of regional trends leaving the gravity reflection of the valley fill. The gravity data and interpreted geologic model for the nine profiles across Hamlin Valley are shown in Figures 4 through 12.

# 4.1 REGIONAL - RESIDUAL SEPARATION

A fundamental step in gravity interpretation is isolation of the portion of the CBA which represents the geologic feature of interest, in this case the relatively low density valley fill. The portion of the CBA which corresponds to this alluvial material is called the 'residual anomaly'.

The residual anomaly was isolated by first estimating the way the CBA field would have appeared if there had been no valley fill present. This estimated field is called the 'regional' gravity. The regional gravity is subtracted from the CBA to produce the residual anomaly. For this study, the regional field was calculated by linear interpolation between the CB' values at bedrock stations on opposite ends of the profiles. Where only one end of a profile was on bedrock, the regional value on the other end was assigned a quantity consistent with the regional trend in the area. This separation technique is

only approximate. Some regional effects may still remain after the subtraction but the error is probably small compared to the large residual anomaly values on these profiles.

The regional field used for each profile is shown together with the CBA in the top portion of Figures 4 through 12. The residual values along each profile are shown by the crosses (x) in the center portions of Figures 4 through 12.

#### 4.2 DENSITY SELECTION

The construction of a geologic model to account for the residual anomaly requires selection of density values representative of the basin fill and of the underlying rock. Since only very generalized density information is available, the geologic interpretation of the gravity data can only be a coarse approximation. Average in situ density of the alluvial fill material was measured between a depth of 100 and 160 feet by six shallow borings in Hamlin Valley. The observed density range for the soil was 2.0 to 2.3 g/cm<sup>3</sup>. These borings were drilled during the Verification Studies of Hamlin Valley (FY 79, FN-TR-27-IV). The larger density value was used in the modeling process to approximate the overall density increase in the alluvium due to compaction with depth (see Grant and West, 1965).

Published values for carbonate rocks typically range between 2.6 and 2.8 g/cm<sup>3</sup>. The Paleozoic carbonate rocks in Nevada are generally reported to be relatively high in density, on the order of 2.8 g/cm<sup>3</sup>. The Nevada volcanic rocks are highly variable in density, ranging between 2.2 and 2.5 g/cm<sup>3</sup>.

The calculated basin depths are very dependent on the density values assigned to the various valley meterials. A one percent change in the average alluvial fill density will result in a five percent change in the calculated fill thickness.

#### 4.3 MODELING

An iterative computer program that calculates the gravitational field for two-dimensional models was used to establish a thickness of alluvium under each profile. The cross-sectional models appear as a set of either 1 or 0.5-km-wide blocks whose tops are at surface elevation and whose bottoms represent the alluvium-bedrock boundary. The elevations at the bottoms of the blocks were adjusted by iterative computation until the computed gravity anomaly for the valley fill differed by less than 1 milligal from the observed residual gravity anomaly. The computed gravity anomaly from the final model is shown as a continuous line in the center portion of Figures 4 through 12. The resulting basin models are shown in the lowest section of Figures 4 through 12. The cross sections have a five times vertical exaggeration so that gentle slopes appear steep.

The gravity survey of Hamlin Valley indicates a structural basin which was formed as a deep graben bounded by steep faults. The basin appears to be several thousand feet deeper in the center, near profiles HV-5 and HV-6, than it is at either end. Actually, it is effectively terminated near profile HV-8. Another basin appears to lie southeast of profile HV-8. Even though the basin is elongate N-S, there appears to be a

component of E-W deformation. The rock outcrop lines strike approximately E-W for about 6 miles on both sides of the valley. On the east side, this trend occurs near profile HV-2 and on the west side it occurs near HV-3. South of profile HV-6, the strike of the rock outcrops is NW-SE. The gravity profiles indicate that the axis of the basin also shifts in direction, or is offset laterally, at several places in the valley. The axis at profile HV-1 (which may actually be the axis of Snake Valley) appears to be approximately 5 miles east of the axis at profile HV-2, and it apparently trends N-S between profiles HV-2 and HV-3. The basin appears to be wider and more complicated beneath profiles HV-4 and HV-5 and the location of the axis is subject to question. However, it seems to be a mile or so west of the axis at HV-3. The orientation of profiles HV-7 and HV-8is such that they probably do not show the axis of the structural basir.

Steep gravity gradients in this basin and range valley are interpreted as being caused by bedrock faults. The faults are interpreted on the gravity profiles where the gradients are maximum. See Figure 13 for an interpretation of possible fault relationships between the profiles. Major-range bounding faults can be interpreted with confidence on profiles HV-2 through HV-6. The displacement associated with the boundary faults increases progressively from profiles HV-2 through HV-5, and decreases from HV-5 to HV-6. If these faults intersect profiles HV-7 and 8, they cross at oblique angles and show only a component of offset.

Neither of the boundary faults show clearly on profile HV-1. According to the trend established from profile HV-5 to HV-2, the offset at HV-1 would be relatively small and the gravity expression would be correspondingly subtle. If the suggestion that there has been E-W deformation between profiles HV-1 and HV-2 is correct, HV-1 may have crossed only the west-boundary fault.

Hintze (1963) shows a fault covered by alluvium along the eastern boundary of the valley. However, the gravity interpretation indicates that there is a pediment-like feature on the east side of the valley and that rock extends westward from the outcrop at shallow depths for 2 to 4 miles (3 to 7 km) before being faulted downward. No corresponding feature is shown on the western valley flanks, where the boundary fault appears to be much nearer the rock outcrop line.

Profiles HV-4 and HV-5 indicate that there may be one or more smaller faults within the basin, forming a local graben along the western boundary. The interpretation, Figure 13, shows a fault roughly parallel to and approximately 1 mile (1.6 km) east of the western boundary fault.

The gravity survey shows the southeastern end of the valley to suddenly become substantially shallower at Profile lines 7 and 8. The maximum depth of the basin shown on these profiles varies from about 1300 to 500 feet (397 to 153 m). The alluvial fill is very thin at the center of each of these profiles suggesting an E-W trending bedrock ridge extending toward the extrusive

volcanic rock outcrops mapped in this area. This bedrock ridge could restrict north-south movement of ground water. An alternative interpretation of profiles HV-7 and 8 is that they cross a shallow flow of igneous material with a density greater than the alluvium. If this is the case there may be substantial thickness of alluvium beneath the flow.

Profile HV-9 crosses the southern end of the valley where it trends NW-SE. The profile extends across the Needle Range into Pine Valley. The pediment feature is absent here, and the gravity interpretation is consistent with the fault suggested by Hintze (1963) in this part of the valley. This portion of the Needle Range is interpreted to be a horst since a fault is interpreted also on the western flank of Pine Valley. A boundary fault is also suggested near the southwestern end of profile HV-9. It is assumed that this fault runs parallel to the northern boundary of the White Rock mountains. It may also be seen on the southern end of profile HV-8. The depth of the basin beneath profile HV-9 is about 2800 feet (853 m).

#### 5.0 CONCLUSIONS

The interpretation of the gravity survey of Hamlin Valley indicates that there are major range bounding normal faults on both sides of the valley. The major graben block between these boundary faults is oriented NNE-SSW. It is calculated to be between 2450 feet (747 m) deep in the north end of the valley and 6700 feet (2042 m) deep near the valley center.

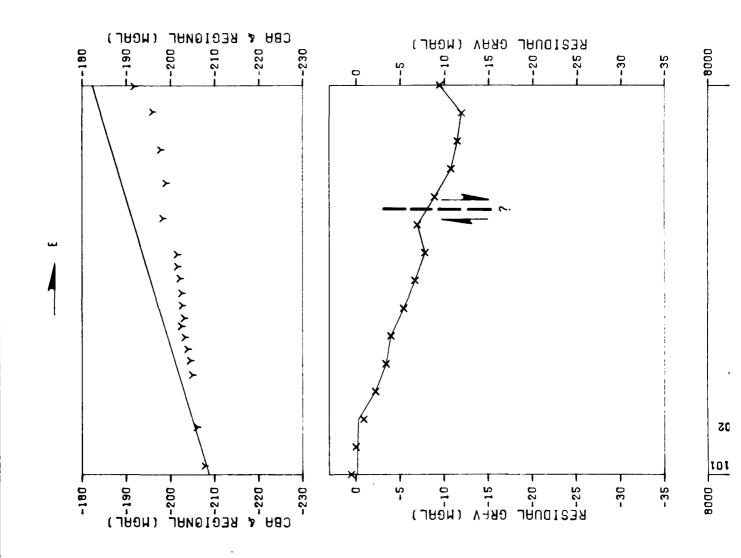
There may be a smaller graben trending NNW along the western side of the major block, near profiles HV-4 and HV-5. This smaller block appears to be from 1000 to 2000 feet (305 to 610 m) deeper than the main block.

A bedrock ridge, which could restrict all but the shallowest ground water movement is interpreted near profiles HV-7 and HV-8.

There is a large, well defined negative gravity anomaly associated with Hamlin Valley. An average density contrast of 0.50 g/cm<sup>3</sup> between the alluvium and bedrock was used to calculate the thickness of the valley fill material. The calculated bedrock depth can only be an approximation since little is known about the actual density distribution in and around the valley. Future studies that acquire better density data or depth to bedrock in deep parts of the valley can be used to refine the gravity interpretation.

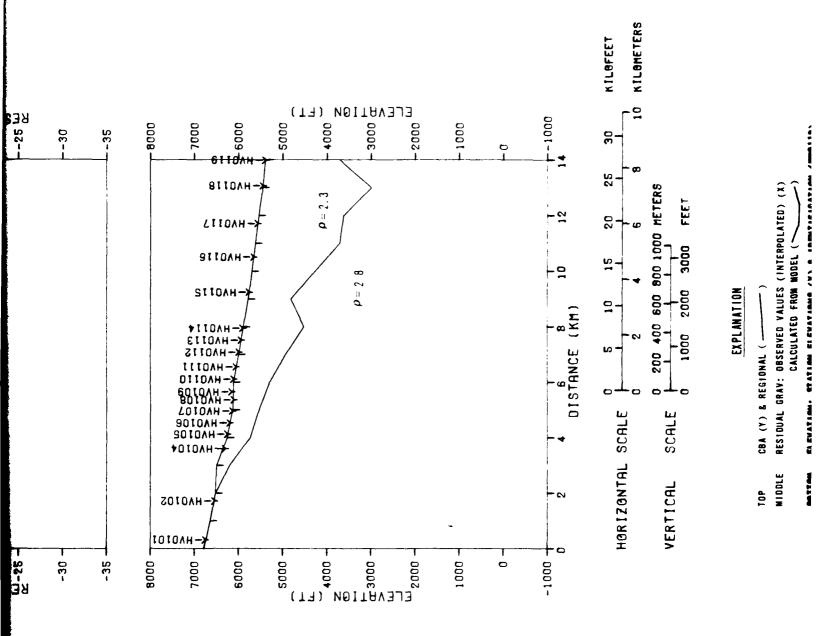
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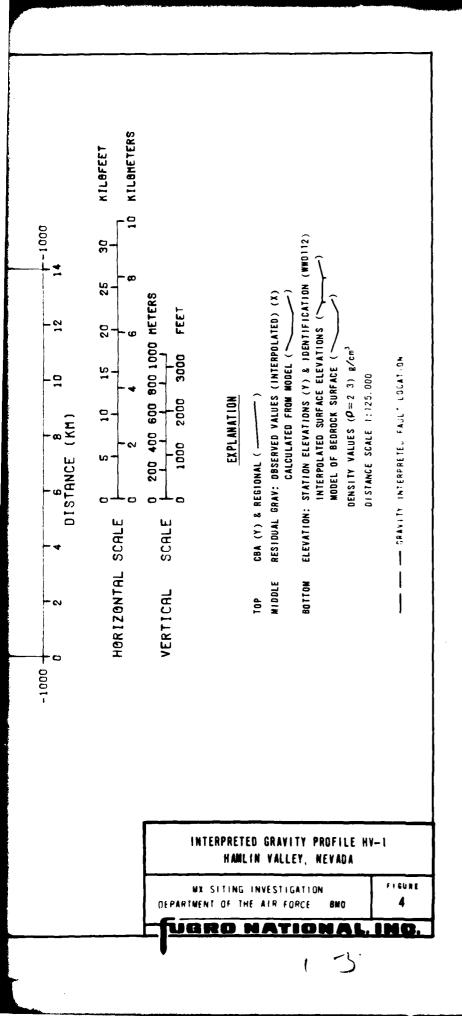
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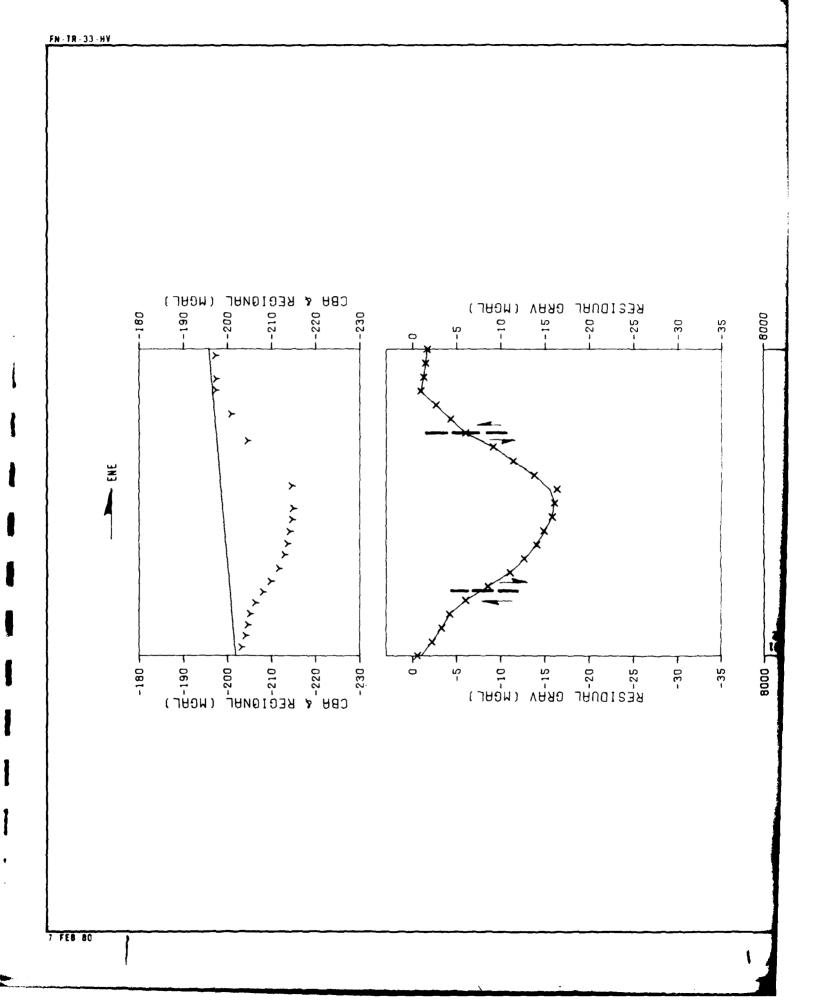


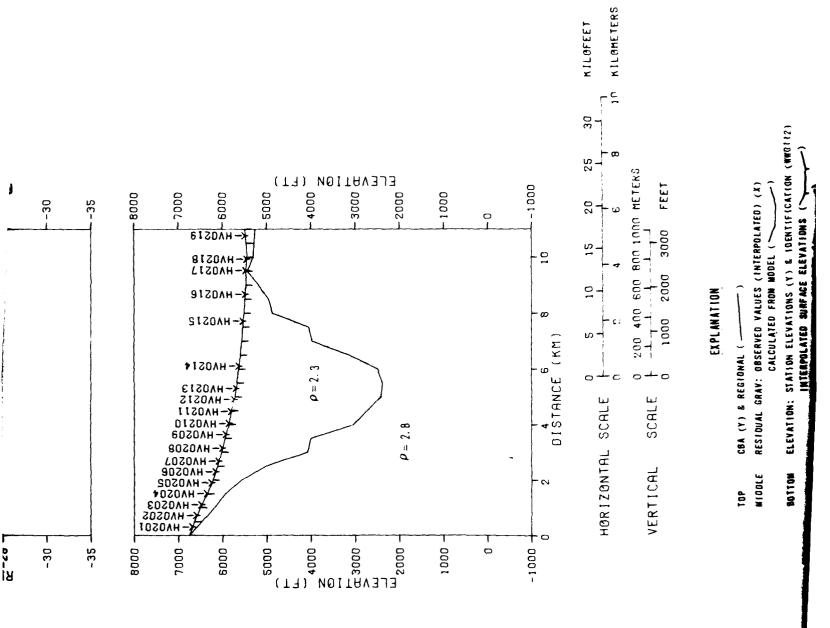
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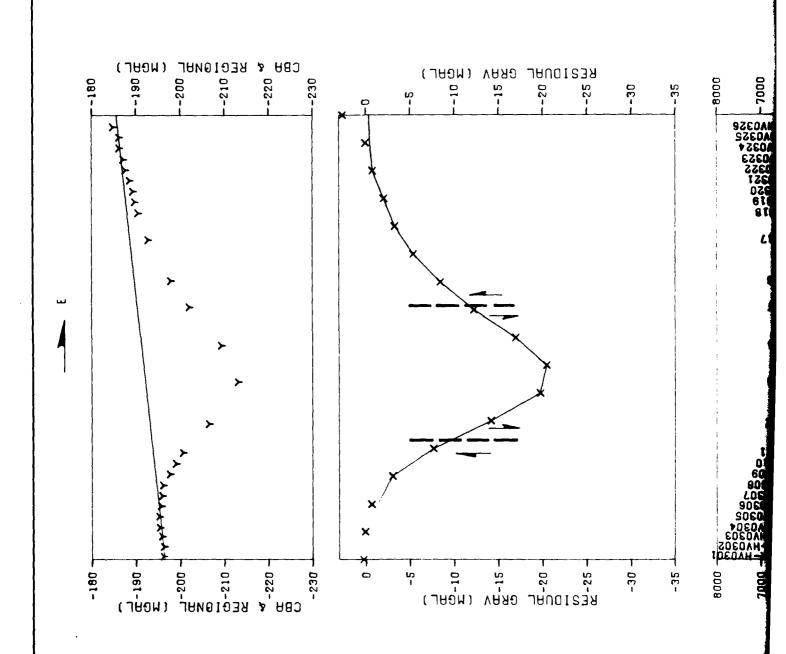
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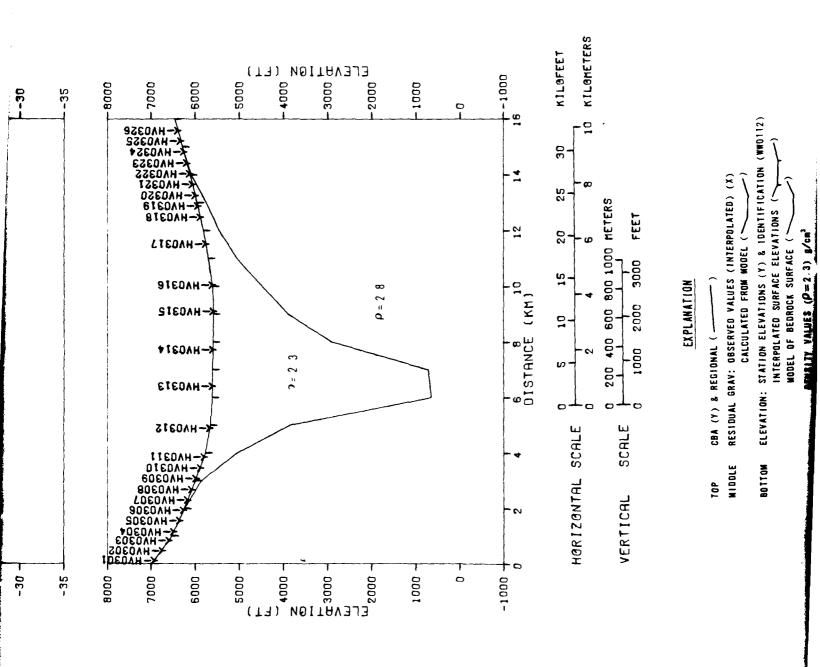
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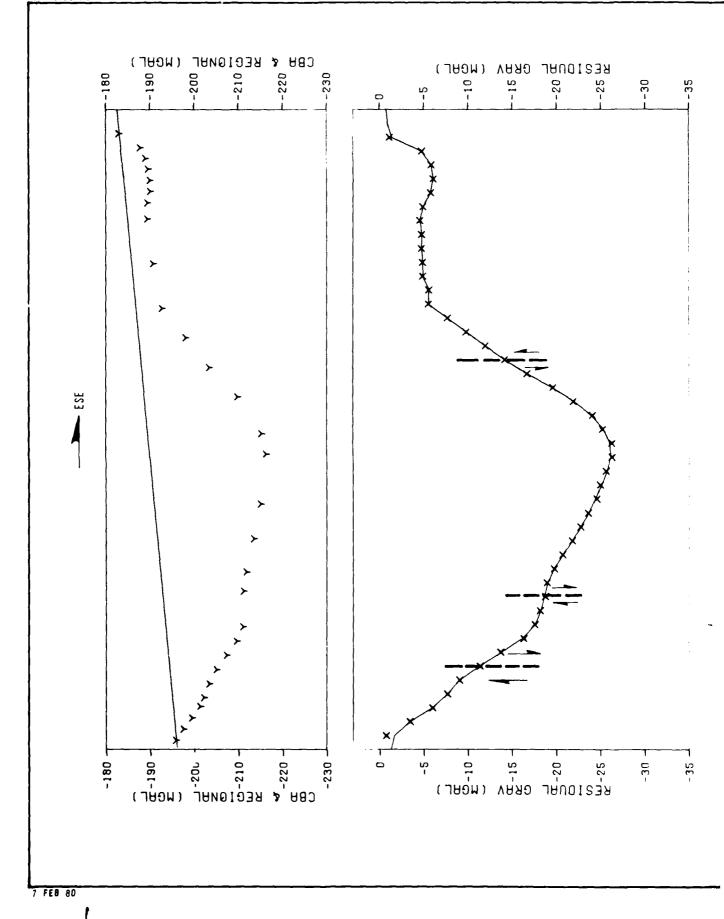
ELEVATION: STATION ELEVATIONS (Y) & IDENTIFICATION (WWO112) RESIDUAL GRAY: OBSERVED VALUES (INTERPOLATED) (X) INTERPOLATED SURFACE ELEVATIONS ( ) 0 200 400 600 800 1000 METERS CALCULATED FROM MODEL ( -GRAVITY INTERPRETED FAULT LOCATION DENSITY VALUES (P=2 3) g/cm3 1000 2000 3000 DISTANCE SCALE 1:125.000 EXPLANATION CBA (Y) & REGIONAL ( --SCALE MIDDLE BOTTOM T0P VERTICAL INTERPRETED GRAVITY PROFILE HV-3

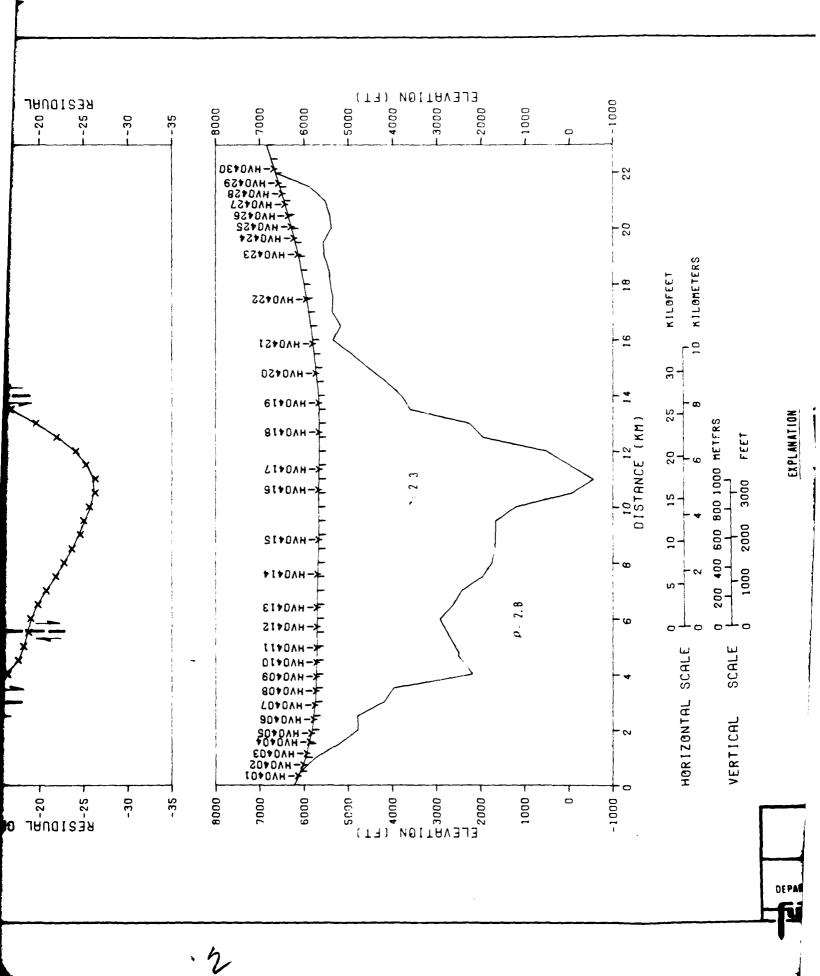
HAMLIN VALLEY NEVADA

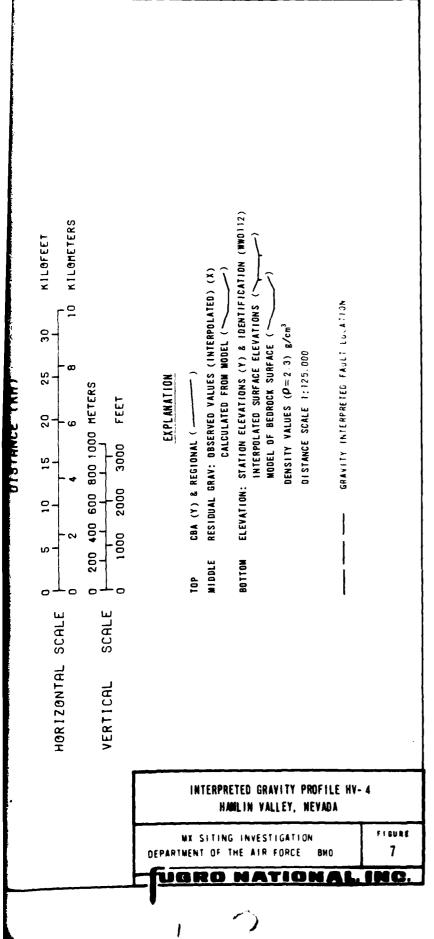
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DEPARTMENT OF THE AIR FORCE

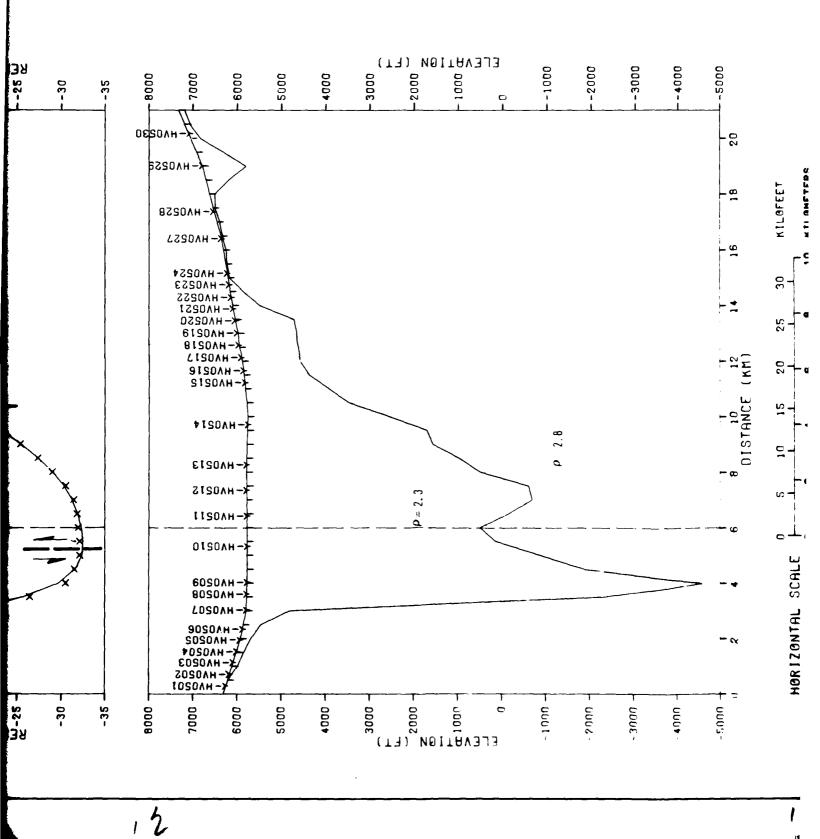
FIGURE





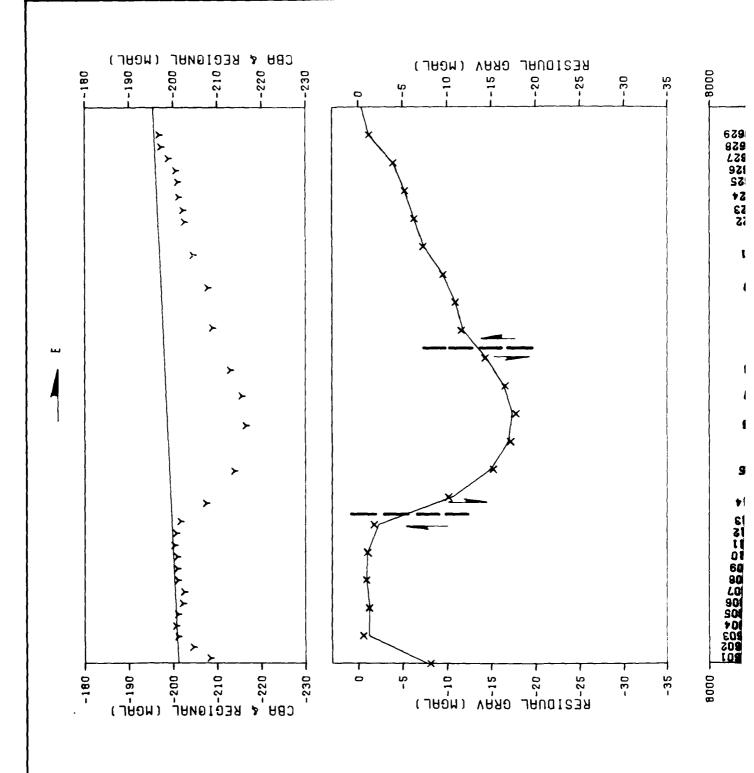


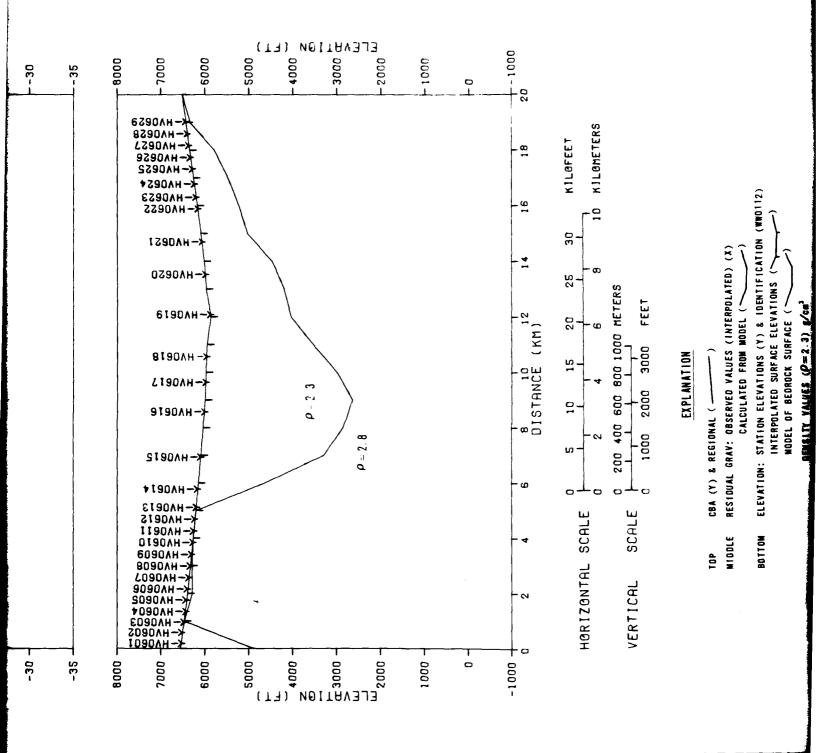
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INTERPRETED GRAVITY PROFILE
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KILOMETERS ELEVATION: STATION ELEVATIONS (Y) & IDENTIFICATION (WWO!12) RESIDUAL GRAY: OBSERVED VALUES (INTERPOLATED) (X) INTERPOLATED SURFACE ELEVATIONS ( 200 400 600 800 1000 HETERS CALCULATED FROM MODEL ( / MODEL OF BEDROCK SURFACE ( -DENSITY VALUES (P=2 3) g/cm3 GRAVITY INTERPRETED FAULT LOCATION DISTANCE SCALE 1:125,000 3000 EXPLANATION 1000 2000 CBA (Y) & REGIONAL ( --HERIZENTAL SCALE SCALE MIDDLE BOTTOM 401 VERTICAL

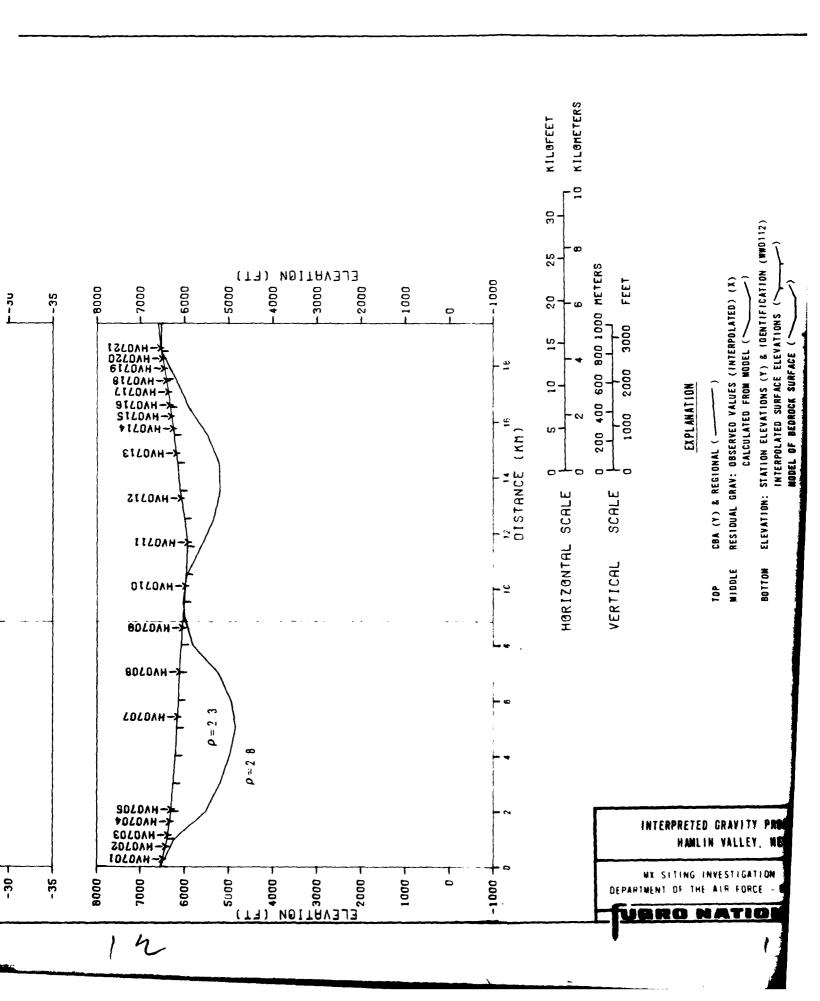
INTERPRETED GRAVITY PROFILE HV-6
HAMLIN VALLEY, NEVADA

WX SITING INVESTIGATION DEPARTMENT OF THE AIR FORCE. SM

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1000 2000 3000 FEET VERTICAL SCALE

EXPLANATION

RESIDUAL GRAV: OBSERVED VALUES (INTERPOLATED) (X)
CALCULATED FROM MODEL ( CBA (Y) & REGIONAL ( --MIDDLE 10P

ELEVATION: STATION ELEVATIONS (Y) & IDENTIFICATION (WWO112) 80TTOM

INTERPOLATED SURFACE ELEVATIONS ()

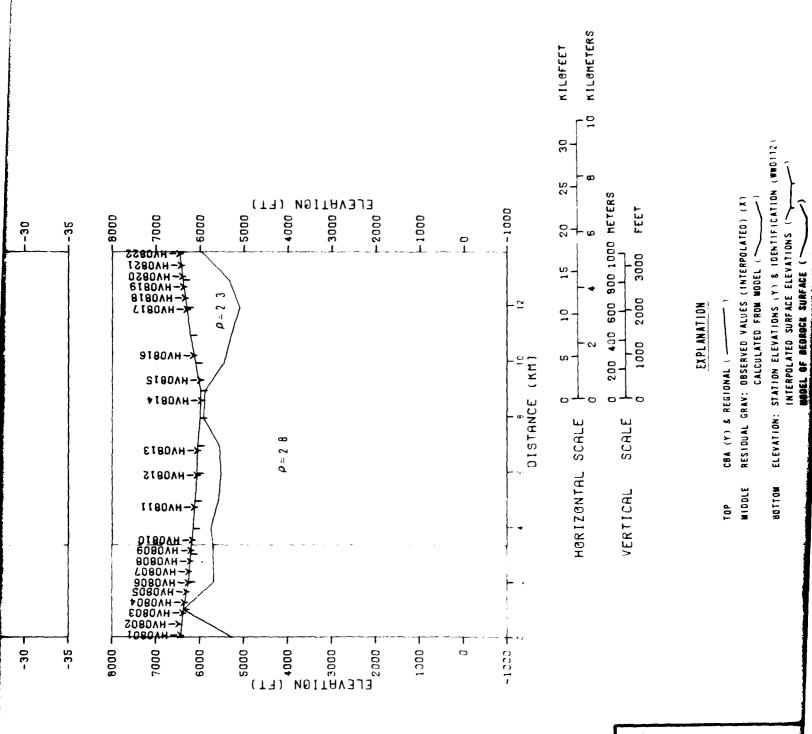
DENSITY VALUES (P=2.3) g/cm3 DISTANCE SCALE 1:125,000 GRAVITY INTERPRETED FAULT LOCATION

INTERPRETED GRAVITY PROFILE HY-7 HAMLIN VALLEY, NEVADA

MX SITING INVESTIGATION

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DEPARTMENT OF THE AIR FORCE



INTERPRETED GRAVITY PM HAMLIN VALLEY, N

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EXPLANATION

CBA (Y) & REGIONAL ( --10P

RESIDUAL GRAY: OBSERVED VALUES (INTERPOLATED) (X) MIDDLE

ELEVATION: STATION ELEVATIONS (Y) & IDENTIFICATION (WWO112) CALCULATED FROM MODEL ( ~ BOTTOM

INTERPOLATED SURFACE ELEVATIONS (>

DENSITY VALUES ( $\rho$ =2.3) g/cm<sup>3</sup>

DISTANCE SCALE 1:125,000

- GRAVITY INTERPRETED FAULT LOCATION

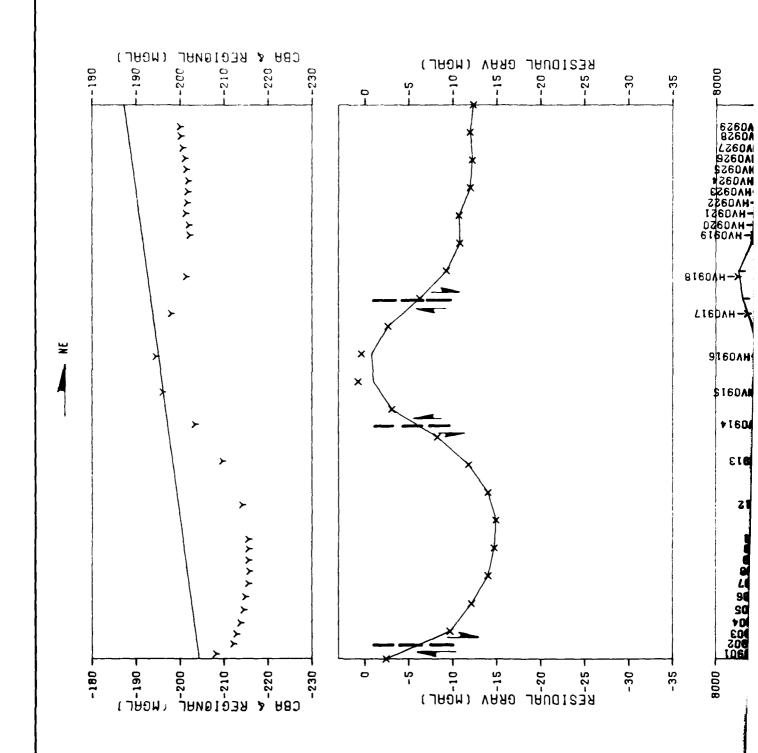
INTERPRETED GRAVITY PROFILE HV-8 HAMLIN VALLEY, NEVADA

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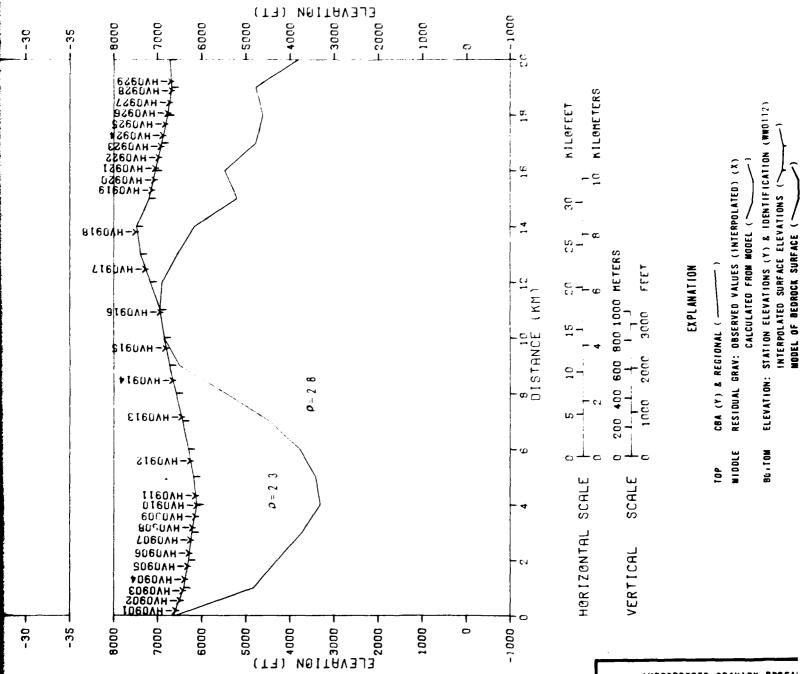
FIGURE

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INTERPRETED GRAVITY PROFIL
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EXPLANATION

CBA (Y) & REGIONAL ( --109

RESIDUAL GRAV: OBSERVED VALUES (INTERPOLATED) (X)
CALCULATED FROM MODEL ( MIDDLE

ELEVATION: STATION ELEVATIONS (Y) & IDENTIFICATION (WW0112) INTERPOLATED SURFACE ELEVATIONS ( BOTTOM

MODEL OF BEDROCK SURFACE ( -

DENSITY VALUES (P=2.3) g/cm3 DISTANCE SCALE 1:125.000

GRAVITY INTERPRETED FAULT LOCATION

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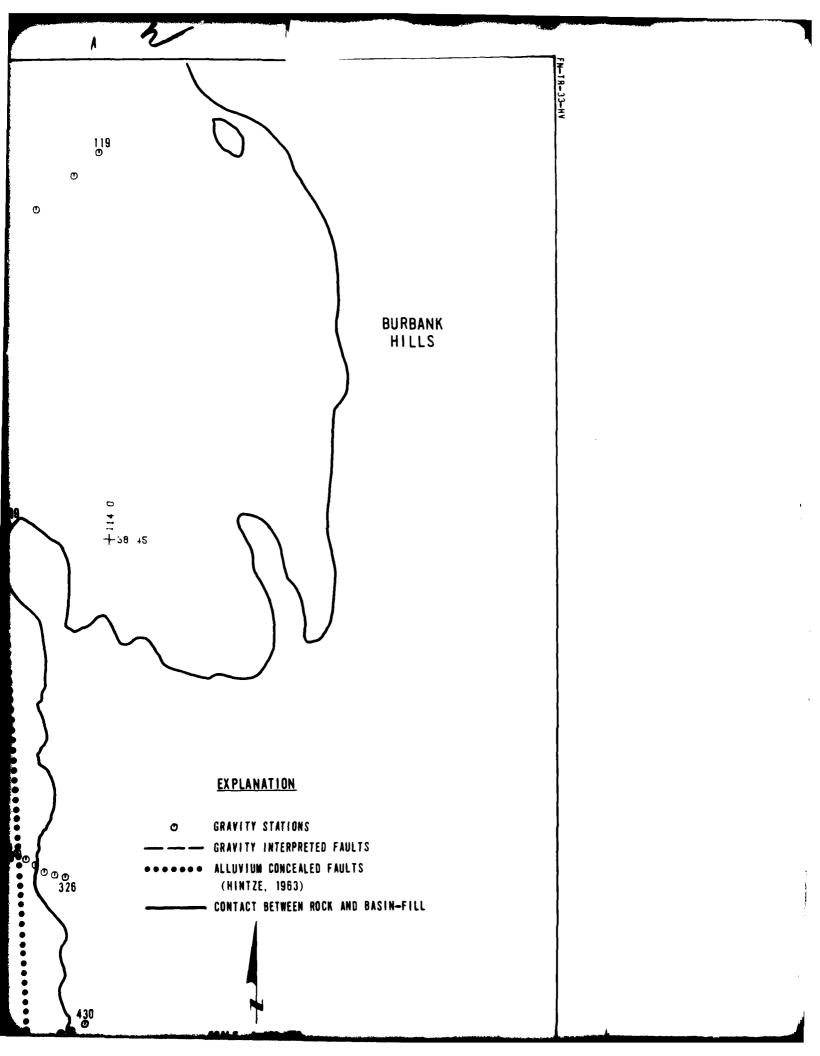
RETED GRAVITY PROFILE HV-9 NAMLIN VALLEY, NEVADA

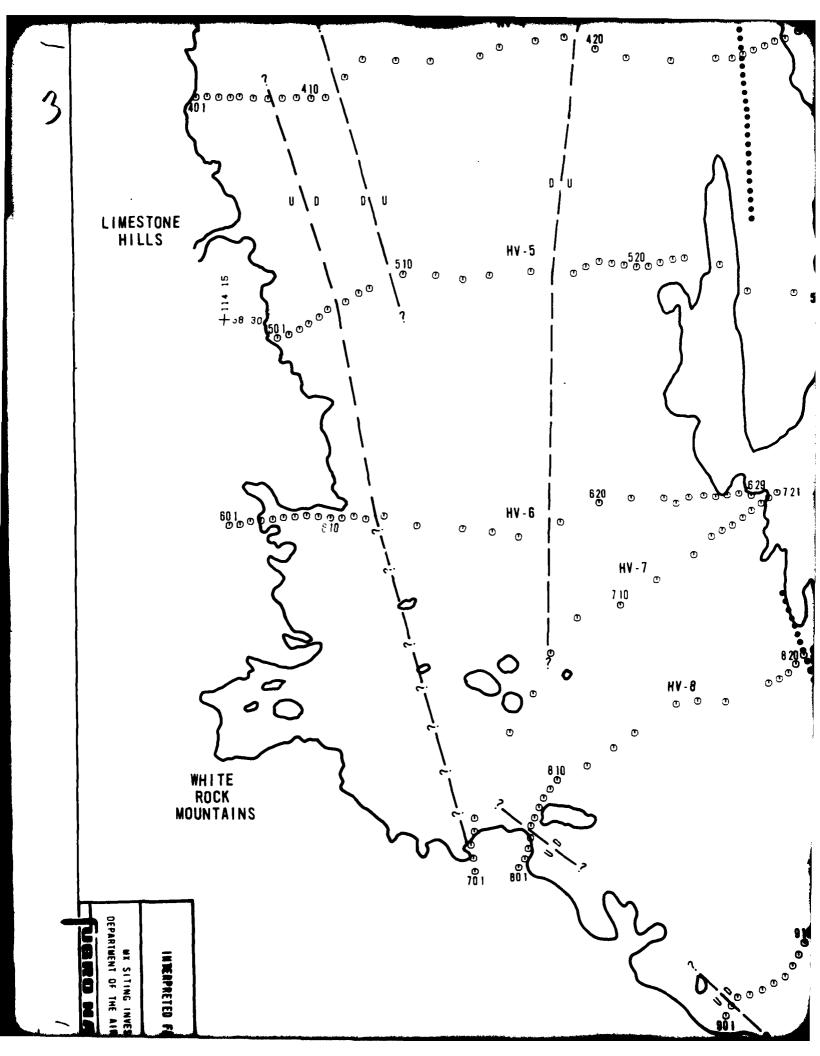
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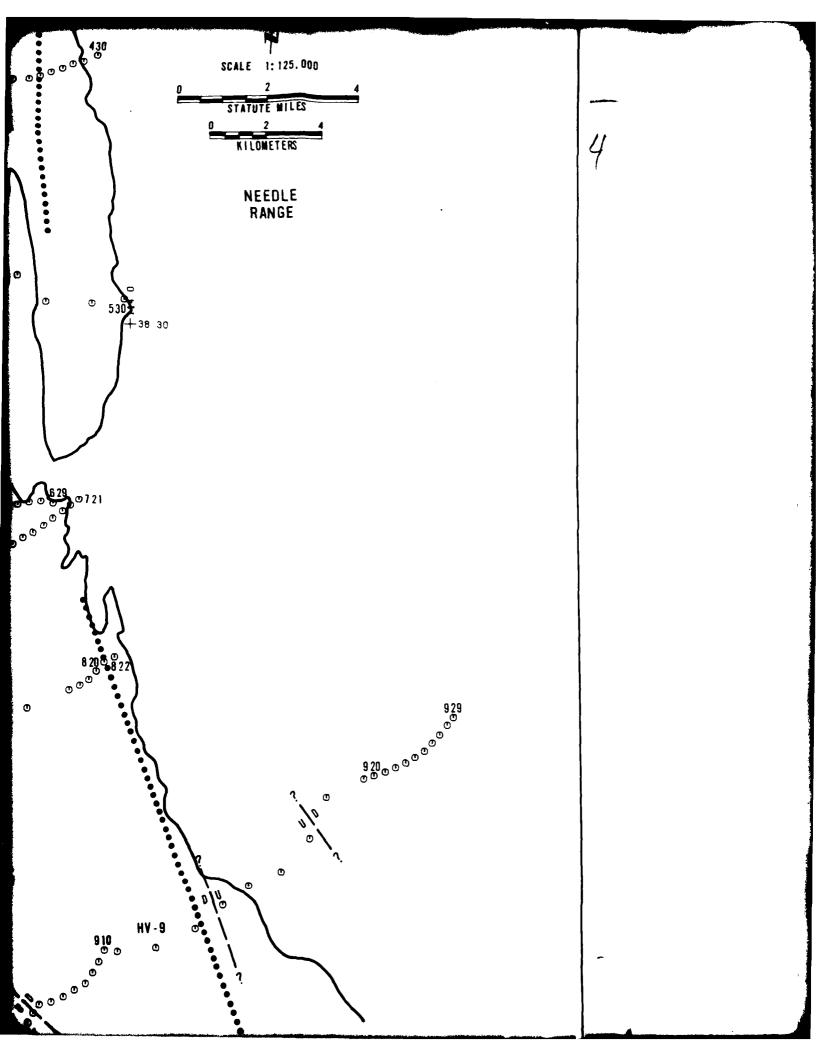
FIGURE

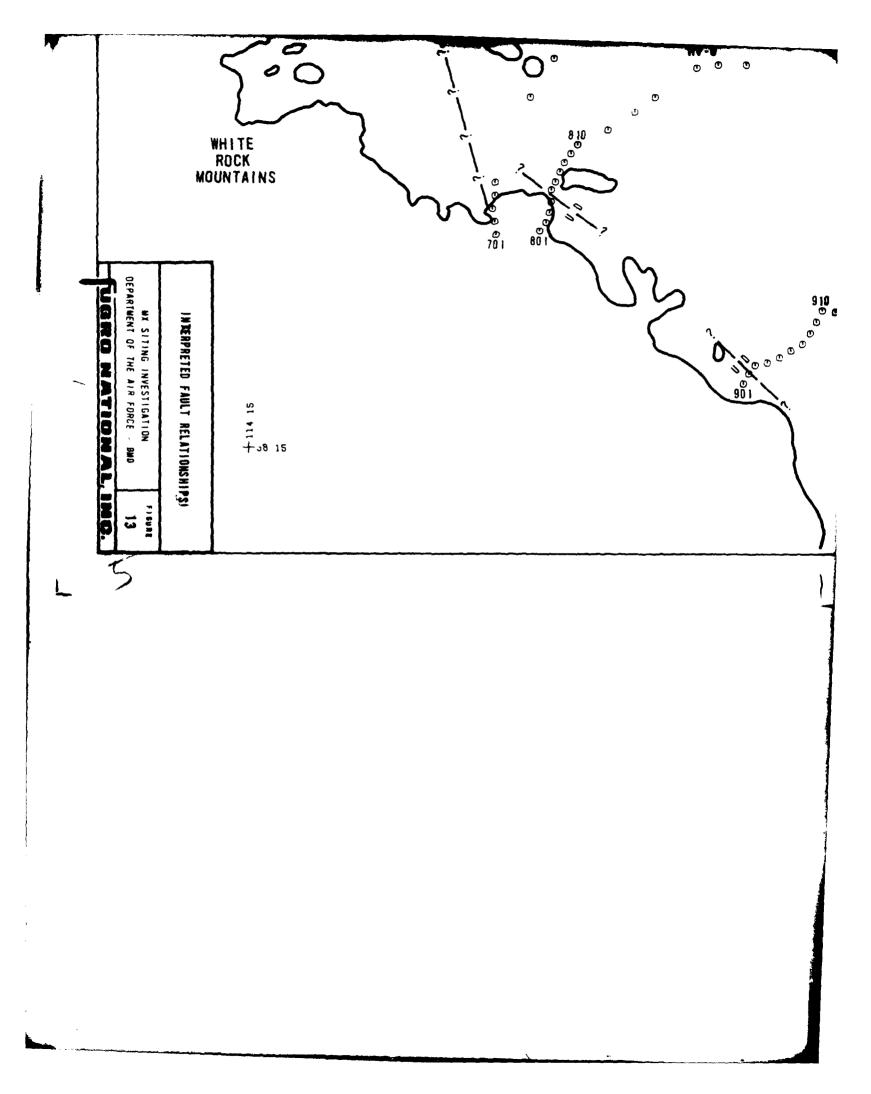
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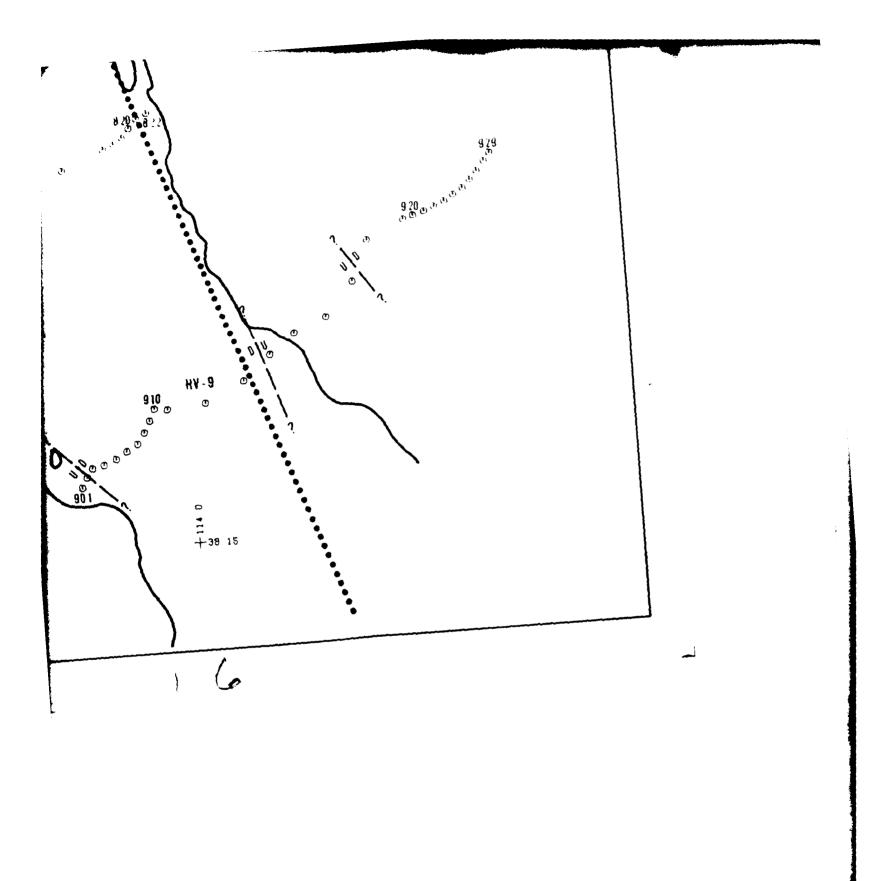
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APPENDIX A1.0

GENERAL PRINCIPLES OF THE GRAVITY EXPLORATION METHOD

# Al.0 GENERAL PRINCIPLES OF THE GRAVITY EXPLORATION METHOD

## A1.1 GENERAL

A gravity survey involves measurement of differences in the gravitational field between various points on the earth's surface. The gravitational field values being measured are the same as those influencing all objects on the surface of the earth. They are generally associated with the force which causes a 1 gm mass to be accelerated at 980 cm/sec<sup>2</sup>. This force is normally referred to as a 1 g force.

Even though in many applications the gravitational field at the earth's surface is assumed to be constant, small but distinguishable differences in gravity occur from point to point. In a gravity survey, the variations are measured in terms of milligals. A milligal is equal to 0.001 cm/second<sup>2</sup> or 0.00000102 g. The differences in gravity are caused by geometrical effects, such as differences in elevation and latitude, and by lateral variations in density within the earth. The lateral density variations are a result of changes in geologic conditions. For measurements at the surface of the earth, the largest factor influencing the pull of gravity is the density of all materials between the center of the earth and the point of measurement.

To detect changes produced by differing geological conditions, it is necessary to detect differences in the gravitational field as small as a few milligals. To recognize changes due to

geological conditions, the measurements are "corrected" to account for changes due to differences in elevation and latitude.

Given this background, the basic concept of the gravitational exploration method, the anomaly, can be introduced. If, instead of being an oblate spheroid characterized by complex density variations, the earth were made up of concentric, homogeneous shells, the gravitational field would be the same at all points on the surface of the earth. The complexities in the earth's shape and material distribution are the reason that the pull of gravity is not the same from place to place. A difference in gravity between two points which is not caused by the effects of known geometrical differences, such as in elevation, latitude, and surrounding terrain, is referred to as an "anomaly."

An anomaly reflects lateral differences in material densities. The gravitational attraction is smaller at a place underlain by relatively low density material than it is at a place underlain by a relatively high density material. The term "negative gravity anomaly" describes a situation in which the pull of gravity within a prescribed area is small compared to the area surrounding it. Low-density alluvial deposits in basins such as those in the Nevada-Utah region produce negative gravity anomalies in relation to the gravity values in the surrounding mountains which are formed by more dense rocks.

The objective of gravity exploration is to deduce the variations in geologic conditions that produce the gravity anomalies identified during a gravity survey.

#### A1.2 INSTRUMENTS

The sensing element of a LaCoste and Romberg gravimeter is a mass suspended by a zero-length spring. Deflections of the mass from a null position are proportional to changes in gravitational attraction. These instruments are sealed and compensated for atmospheric pressure changes. They are maintained at a constant temperature by an internal heater element and thermostat. The absolute value of gravity is not measured directly by a gravimeter. It measures relative values of gravity between one point and the next. Gravitational differences as small as 0.01 milligal can be measured.

# A1.3 FIELD PROCEDURES

The gravimeter readings were calibrated in terms of absolute gravity by taking readings twice daily at nearby USGS gravity base stations. Gravimeter readings fluctuate because of small time-related deviations due to the effect of earth tides and instrument drift. Field readings were corrected to account for these deviations. The magnitude of the tidal correction was calculated using an equation suggested by Goguel (1954):

 $C = P + N\cos \phi (\cos \phi + \sin \phi) + S\cos \phi (\cos \phi - \sin \phi)$  where C is the tidal correction factor, P, N, and S are time-related variables, and  $\phi$  is the latitude of the observation point. Tables giving the values of P, N, and S are published annually by the European Association of Exploration Geophysicists.

The meter drift correction was based on readings taken at a designated base station at the start and end of each day. Any difference between these two readings after they were corrected for tidal effects was considered to have been the result of instrumental drift. It was assumed that this drift occurred at a uniform rate between the two readings. Corrections for drift were typically only a few hundredths of a milligal. Readings corrected for tidal effects and instrumental drift represented the observed gravity at each station. The observed gravity values represent the total gravitational pull of the entire earth at the measurement stations.

### A1.4 DATA REDUCTION

Several corrections or reductions are made to the observed gravity to isolate the portion of the gravitational pull which is due to the crustal and near-surface materials. The gravity remaining after these reductions is called the "Bouquer Anomaly." Bouquer Anomaly values are the basis for geologic interpretation. To obtain the Bouquer Anomaly, the observed gravity is adjusted to the value it would have had if it had been measured at the geoid, a theoretically defined surface which approximates the surface of mean sea level. The difference between the "adjusted" observed gravity and the gravity at the geoid calculated for a theoretically homogeneous earth is the Bouquer Anomaly.

Four separate reductions, to account for four geometrical effects, are made to the observed gravity at each station to arrive at its Bouquer Anomaly value.

a. Free-Air Effect: Gravitational attraction varies inversely as the square of the distance from the center of the earth. Thus corrections must be applied for elevation. Observed gravity levels are corrected for elevation using the normal vertical gradient of:

FA = -0.09406 mg/ft (-0.3086 milligals/meter) where FA is the free-air effect (the rate of change of gravity with distance from the center of the earth). The free-air correction is positive in sign since the correction is opposite the effect.

b. Bouguer Effect: Like the free-air effect, the Bouguer effect is a function of the elevation of the station, but it considers the influence of a slab of earth materials between the observation point on the surface of the earth and the corresponding point on the geoid (sea level). Normal practice, which is to assume that the density of the slab is 2.67 grams per cubic centimeter was followed in these studies. The Bouguer correction ( $B_c$ ), which is opposite in sign to the free-air correction, was defined according to the following formula.

 $B_{c} = 0.01276 (2.67) h_{f} (milligals per foot)$ 

 $B_C = 0.04185$  (2.67)  $h_m$  (milligals per meter)

where  $\mathbf{h}_{\mathbf{f}}$  is the height above sea level in feet and  $\mathbf{h}_{m}$  is the height in meters.

c. Latitude Effect: Points at different latitudes will have different "gravities" for two reasons. The earth (and the geoid) is spheroidal, or flattened at the poles. Since points at higher latitudes are closer to the center of the earth than points near the equator, the gravity at the higher latitudes is larger. As the earth spins, the centrifugal acceleration causes a slight decrease in gravity. At the higher latitudes where the radii of the circles of latitude are smaller, the centrifugal acceleration diminishes. The gravity formula for the Geodetic Reference System, 1967, gives the theoretical value of gravity at the geoid as a function of latitude. It is:

g = 978.0381 (1 + 0.0053204  $\sin^2 \emptyset$  - 0.0000058  $\sin^2 2\emptyset$ ) gals where g is the theoretical acceleration of gravity and  $\emptyset$  is the latitude in degrees. The positive term accounts for the spheroidal shape of the earth. The negative term adjusts for the centrifugal acceleration.

The previous two corrections (free air and Bouguer) have adjusted the observed gravity to the value it would have had at the geoid (sea level). The theoretical value at the geoid for the latitude of the station is then subtracted from the adjusted observed gravity. The remainder is called the Simple Bouguer Anomaly (SBA). Most of this gravity represents the effect of material beneath the station, but part of it may be due to irregularities in terrain (upper part of the Bouguer slab) away from the station.

d. Terrain Effect: Topographic relief around the station has a negative effect on the gravitational force at the station. A nearby hill has upward gravitational pull and a nearby valley contributes less downward attraction than a nearby material would have. Therefore, the corrections are always positive. Corrections are made to the SBA when the terrain effects were 0.1 milligal or larger. Terrain corrected Bouguer values are called the Complete Bouguer Anomaly (CBA). When the CBA is obtained, the reduction of gravity at individual measurement points (stations) is complete.

## A1.5 INTERPRETATION

The first step in interpretation is to separate the portion of the CBA that might be caused by the lightweight, basin-fill material overlying the heavier bedrock material which forms the surrounding mountains and presumably the basin floor. Since the valley-fill sediments are absent at the stations read in the mountains, the CBA values at these bedrock stations are used as the basis for constructing a regional field over the valley. A regional field is an estimation of the values the CBA would have had if the light weight sediments (the anomaly) had not been there.

The difference between the CBA and the regional field is called the "residual" field or residual anomaly. The residual field is the interpreter's estimation of the gravitational effect of the geologic anomaly. The zero value of the residual anomaly is not exactly at the rock outcrop line but at some

distance on the "rock" side of the contact. The reason for this is found in the explanation of the terrain effect. There is a component of gravitational attraction from material which is not directly beneath a point.

If the "regional" is well chosen, the magnitude of the residual anomaly is a function of the thickness of the anomalous (fill) material and the density contrast. The density contrast is the difference in density between the alluvial and bedrock material. If this contrast were known, an accurate calculation of the thickness could be made. In most cases, the densities are not well known and they also vary within the study area. In these cases, it is necessary to use typical densities for materials similar to those in the study area.

If the selected average density contrast is smaller than the actual density contrast, the computed depth to bedrock will be greater than the actual depth and vice-versa. The computed depth is inversely proportional to the density contrast. A ten percent error in density contrast produces a ten percent error in computed depth. An iterative computer program is used to calculate a subsurface model which will yield a gravitational field to match (approximately) the residual gravity anomaly.

APPENDIX A2.0

HAMLIN VALLEY

GRAVITY DATA

## PROFILE #1 HAMITM VALUEY GRAVITY DATA

IDENT.	LAT. DEG MIN	DEG	MIN	+0006	7 N	CUI	1,15	ti I 🥶	GPAV	GRAV		CBA 41000
HV0101	385147	114	962	6733H	3	470	430463	74041	145080	206/63		79205
Hvnin2	385159									206780	1212	79377
HV0103	385155	114	749	68638				-		206774	1708	79451
HVOlcu	365147	114	735	6286v	t.	325	430473	74969	144) 95	206763	597	794A1
HV0105	385143	114	699	62376	0	281	130467	75022	144581	200750	520	79530
HV0106	385138	114	671	61030	V!	2886	150449	75064	148951	205749	ج به ت	79592
HV0107	385128	114	641	61266	U	2714	1301"2	75107	149393	206734	217	79653
HV0108	385123	114	513	60955	Ü	2580	130434	75147	149024	206727	265	79733
HV0109	385100	114	593	61418	Ü	2720	430403	75177	149299	206702	ے د، <u>د</u>	79670
HV0110	385109	114	561	60046	Ų	2090	112051	75223	144037	206/00	2.95	79709
HVn111	345112	114	531	60434	(.	1941	356417	15261	LAGURG	20e/11	134	19728
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HV0113	385113	114	465	5929V	C	1820	430422	75362	150775	206713	-130	19020
HV0114	385109	112	435	54424	f)	1746	13041b	75400	151008	206706	-245	79032
HV0115	385105	114	345	5751H			130415			206702	<b>≈</b> 335	50156
HV0116	385120	114	258	5651 v	· ·	1446	430445	75661	152753	206723	-791	00051
HV0117	385140	114	175	55634	$t_{J}$	130	450 165	75780	153440	260752		80203
HVOLIA	365205	114	A 1	54395	ij	1.30/	436611	15912	154465	205043	-1194	60391
HV0113	385250	114	18	53778					155310			808 <b>1</b> 0

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PROFILE #2 HAMIN VALLEY GRAVITY DATA

STATION TOENT.	LAT. DEG MIN	DEC	MIN	+CCIUE	TO	(QUT	UTM	Lilw	GRAV	THEC	FAA	+1000
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HV0201	384805						124834			266259	1563	79574
HAUSUS	364795						12961/			206245	1179	79521
HV0203	384787				•		129203			206233	714	79472
HAU504	384784				•	. •	120744			855005	291	79354
HV0505	384783						129798			155805	-175	7918c
HV0506	384775	1 1 4	777	- ウェンコン	•	_	429785			206215	-593	79007
HV0207	384772					201	429796			206210	-1932	78617
HA0508	384763					274	429765	75082	14912	1200197	-1419	78686
HV0504	384759						120750			206191	~1085	78014
HV0510	384745									206170	<b>≈1900</b>	78572
FV0511							429708			3200148	-2159	78514
HVOSIS	384730				•		950689		150261	206132	-2320	78496
HV0213	384719						429670			5510055		
HV0214	384712						429650		:152312	12/5100	-1784	19527
HV0215	384697			5525 v			429638	7557	115046	აგიგიგა	-1562	79699
HV0216	384686				•	•	429574	75620	315 437	1206030	-1205	80295
HV0517	384650	114	243	79513	, ,,		420459		115340	32115941	-1344	60230
HA0518	384588						467437 420359	7560	4 L - 3 T O . 7 1 5 Z O O !	5205852	-11A	B 1254
HV0519	384528	114	529	.) (1 14 (1);	, ()	161	116.1.25.1	1 30 4	112341	7 to 10 10 10 10 10		

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PROFILE #3 HAMITY VALLEY CRAVITY DATA

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HV0302	384212	1141046			•				205360		80397
HV0303	384194	1141659				28690 28660			205335		80446
HV0304	384177	1141014		•		28550 28550			205322		69463
HV0305	384168	114 989				28627			205302		80423
HV0306	384155	114 968			•	28595			205276		80396
HV0307	384137	114 951		•	-				205259		80375
HV0308	384125	114 929		-		28573			205248		80221
HV0309	384118	114 902				38562			205247		80088
HV0310	384117	114 874			-	28561	•		245241		79927
HV0311		114 847				24555			1205241		79355
212 nvn		114 767			•	24460				-2340	78677
HV0313	-	114 040		=	•	23477			1205173	-1980	79(55
HVn314	384619	114 608				322			205102		79783
HV0315			5581Y	-		158317			\$26503d	٠١٥١٠ حـ -	30197
HV0316	383972	• .			•	148 81 5			<b>.</b> 205034	218	51708
FV0317		114 360			-	128261			1500000	<del>-</del>	•
HV0318	383415		58778			15451			7494054	1 6 K	#0938
HV0319	383906	• •		-	•	144139			12,4937	1(88	81009
HV0320	383898	114 259			•	158184			2504955	1314	81049
HV0321	383886				1714	154161			7204910	1613	
HV0322	383875				1836	158100			1649000		81227
FV0323	383864	114 191			•	128125			4204675	5100	81274
HV0324	383850	114 169	67486	v 0	204/	158100			0204054	2471	81364
HV0325	383844	114 14	63238	i u	2100	158960			2704845	2714	81367
HV0326	383840	114 117	43865	, 1	252	158084	75442	147/7	3594434	303n	9150B

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PROFILE AGENCITY DATA

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HV0401 383440	• • • •	8 303427278	73665147553204252	1015 80411
HV0402 383441	1141536 6012B	0 152427281	73905148260204253	
HV0403 383439		0 145427279	73946148531204250	
HV0404 383438	• •	0 135427278	73988148805204246	
HV0405 383439	• •	0 130427281	74020148940204250	
HV0406 383436		u 126427277	74073149112204245	<b>~</b> 775 79652
HV0407 383435		0 113427277	70155100150504502	
HV0408 383436		0 115427280	7/1173148995204245	
HV0409 383437		6 110427283		
HV0410 383436	• • • • • • • • • • • • • • • • • • • •	0 106427283		
HV0411 383437	•	0 104427267	74328148630204247	
HV0415 383476	• • • •	0 102427361	74396148820204305	
HV0413 343511		0 99427428	•	
HV0414 383508	•	0 96427426		
HV0415 38350/		0 98427426		
HV0416 363517		0 98427455		
HV0417 383533		0 102427484	· · · · ·	
HV0418 383546	=	0 164027512	•	
HV0419 383552		0 104427520		
HV0420 383529		0 111427487		
HV0421 383511	• •	0 126427457		
HV0422 383504	-	0 139427449	• • • • •	
HV0453 343200		0 165427458	,	-
HV0424 383504	· •	0 180427460	• • • • • • • • • • • • • • • • • • • •	
HV0425 363513	-	0 189427474		
HV0426 383521		0 213427490	· · · · · · · · · · · · · · · · · · ·	
HV0427 383528		0 216427505		• •
HV0428 383537		0 231427523		
HV0429 383542	•	0.253427533		
HV0430 383550	114 70 6678Y	3 280427560	76028145772204422	4146 41445

FNN UF LIST

PROFILE #5
HAMLIN VALUEY GRAVITY DATA

STATION LAT.	LONG. ELEV. DEG MIN +CODE	TERMOOR. NORTH	EAST DRSV THEC	FAA CBA +1000
* * * * * * * * * * * * * * * * * * * *	. Paris Produc			71000
	1141369 62486		74175146183203552	1440 80406
HV0502 382970	1141340 6163V		74216146599203561	1041 80232
HV0503 302980	1141314 6078V	1 184425435	74254146902203576	527 79983
HV0504 382992	1141291 6000H	0 176426461	74286147349203594	226 79937
HV0505 383004	1141266 59228	0 164426485	74322147663263611	-214 79752
HV0506 383018	1141245 5862Y	0 152426512	74352147887203632	•574 79584
HV0507 383033	1141200 5779Y	0 137426541	74416147580203654	-1685 78742
HV0508 383050	1141167 57696	0 124426574	74463146928203679	<b>-2458 77990</b>
HV0509 383059	1141141 5763Y	0 118426592	74501146663203692	-2700 77672
HV0510 383085	1141057 5759Y	n 100456944	74621146645203730	-2885 77574
HV0511 383083	114 977 5767Y	त । १००४८ ५५४४	74738146683203727	-2150 77662
_HV0512	114 910 57748	0 99420032	74836145761203/16	<b>-</b> 2613 77793
HV0513 383082	114 844 57/91	u 994255 16	74931147017203726	•
HV0514 383089	114 742 57444	0 107426665	75079147915203736	-1762 78754
HV0515 363035	114 636581401	L 110426663	75233148362203730	⇒650 7963)
_HV0516 _ 3d3097	114 506 58518	u 113426586	75276148385203746	-295 79865
	114 573 59014	6 116466700	75323148284203763	54 60013
HV0518 383108	114 542 59506	0.123426694	75369148011205755	351 80145
HV0519 383100	114 512 59959	u 12342n69n	75413147811263752	479 80155
HV0520 383096	114 481 60394	0 133426690	75458147632203746	719 80256
HV0521 383096	114 451 60858	0 137426692	75501147588203746	1107 80491
- HV0522 - 383103	114 422 m130H	0 144156736	75543147467243756	1415 HOF 44
HV0523 384110	114 391 61764	0 147426720	75588147462203767	1755 80901
HV0524 383112	114 362 62196	0 153426725	75030147479203776	2539 81181
HV0527 383098	114 275 6350Y	0 0426764	75757146930203749	2546 61268
HV0528 383046	• • • • • • • • • • • • • • • • • • • •		75859145597203673	3575 81292
HV0529 383042	• •		76025143557203667	3714 80827
HV0530 383050	114 13 70679	0 323426527	76141142208203679	5046 61255

END OF LIST

PROFILE #6
HAMETA VALLEY GRAVITY DATA

STATION LAT. THENT, DEG MIN	LONG. ELEV		OR. NORTH	FAST DRSV		FAA	CBA +1000
	~-~	<u>-</u>			~ _ ~~ ~~ ~~		
HV0601 382595	1141493 652	oc v	159425718	7401514287	0203011		74141
HV0602 382596	1141466 651	SC 0	153425721	7405414331	4203013	1592	79535
HV0603 382602	- •		150425733	7409114395	0203021	1749	79665
HV0604 382604	1141414 641	6H ()	161425738	7412914427	6203v24	1540	77917
HV0605 382000	1191385 641	<b>7</b> 3 0	152425743	7417114424	1203027	15,09	79675
- hv060m - 382009	1141358 638	6년 0	153425750	7421014430	3203031	1437	79100
HV0607 382610	1141330 635	OY U	159425753	7425114450	2203033	1234	79/35
HV0608 382613	1141301 632	0 B	173425760	7429314481	0203038	1250	79873
HV0609 382611	1141272 624	0.3	173425757	7433514500	4203034	1170	79896
HV0610 382608	1141242 625	7B 0	159425753	7437914521	A203030	1075	79894
HV0611 382607	1141214 524	48 0	145425753	7442014537	1203028	1100	74956
HV0615 385010	1141184 622	OB U	130425760	7446314548	5203033	40%	70014
HV0613 382606	1141155 bib	Z Y U	135425755	7450614500	4503057	172	79616
HV0614 382611	1141110 015	7ド ()	12/425765	7457114519	1203034	100	79235
HV0615 302592	1141029 608	4Y U	117425733	7469014497	620300p	-774	78594
- HV0616 - 382586	114 916 599	74 0	109455727	7485514523	2202494	-1321	78329
HV0617 382578	114 842 547	1 Y ()	10/02/5710	7496314547	985489	-1320	78425
- MV061H - 382569	114 777 599	SY ii	100025702	750581 -581	n202913	-1113	74086
HV0619 382596	114 674 586	e.S U	112425757	7520614678	8203013	-1010	79089
HV0620 382634	114 577 597	413 O	111425831	7534514630	6203068	-534	79197
HV0621 352643	114 497 605	GY ()	122425852	7546114613	5203081	79	79536
HV0623 382645	114 416 615	O 45	120425854	7557914576	5203080	584	79727
HV0623 382632	114 386 619	7:- u	120125636	7502316551	0203065	168	79750
- HV0624 382664	114 354 623	66 0	131465860	7506913542	2803083	471	14855
HV0625 382641	114 317 627	5 Y (1	133425857	7572314517	7203087	1149	79680
HV0626 382644	114 288 632	0 40	140425565	7576514494	2203083	1338	79924
HV0527 302649	114 259 639	71: G	142425872	7580714468	52030BB	1632	H0000
HV0628 382651	114 230 639	36 0	148025879	7564914463	9203093	1913	HJ257
HV0629 352646	114 200 546	<b>7</b> m = 0	151425871	7539314466	0203085	2070	89588

FND OF LIST

PROFILE #7
HAMLIN VALLEY GRAVITY DATA

STATION	LAT.	UÜ	VG. I	ELEV.	TEP-0	fiR.	MORTH	FAST	ORSV	THEC	FAA	CBA
	DEG MIN								GRAV	GRAV		+1000
~~~~												
HV0701	381915	114	892	64888	Ü	218	424487	74928	141801	202013	540	78939
HV0702	381943	114	395	64206	6	217	424539	74922	142281	202055	650	78976
Hv0703	381970	114	900	434BH	Ģ	182	424580	74913	142527	202094	573	78834
HV0704	381996	114	893	63298	Ú	165	424637	71472	142716	551509	151	78731
HV0705	55058	114	208	62988	Ų	153	424685	74422	142488	2.,2170	<b>→</b> r,	Inno
HVATAT	382133	114	803	01500	$\mathbf{e}$	130	481390	75042	144171	د 1 نه نم ن <b>ن</b> م	- 35%	76785
HV0708	302863	114	744	61066	()	114	425137	75124	145041	202525	-19	79271
HV0709	382341	114	700	6045H			425284				494	79497
HV0710	382410	114	634	5969Y	ij	110	425414	75275	146926	202739	564	60116
<b>570711</b>	382435	114	527	5919Y	U	154	425465	15430	147044	202776	<b>-</b> 25	79511
HV0712	382483	114	437	p079B	Û	151	425566	75598	145900	505846	27K	19665
HV0713	382531	114	344	0178Y	ij.	155	425652	75590	145441	202917	€76	79721
HY0714	382566	114	300	6255Y	U	131	465718	75/52	145188	202968	1091	79686
HV0715	382578	13	274	62948	J	133	425742	75/89	145091	202966	1540	50008
HV0716	362588	114	249	63258	O	145	425761	75025	145010	203001	1501	80113
HV0717	382602	114	223	6360B	Ü	140	425789	75002	144911	203021	1/40	90197
HV071H	362010	114	201	6400R	Ó	145	425815	15893	144742	203041	19.57	HU253
HV0719	-382030	114	177	n455Y	()	153	425842	75428	148500	203003	5196	A0335
HV0720	382641	114	157	644 देत	(;	150	425564	75450	1.14390	263670	2430	86441
HV0721	382053	114	136	6535Y	0	164	425887				2571	80446

FROM DE LIST

PROFILE AR HAMITA VALLEY GRAVITY DATA

STATION IDENT.	LAT. DEG MIN						NORTH		U35V 624V	THEC	FAA	CBA +1000
HVD801	381923						424507			202025		75492
SUBUNH	381941	114	769	64758	0	192	424541			202052		79074
HV0803	381962	114	760	63820	Ü	166	424580	75115	142722	202082	707	79105
HV0804	381983	114	754	63480	U	154	4246.19	75125	142429	202113	563	79000
HV0805	382007	112	153	62948	G	140	484665	75125	1 4 3 1 9 5	202146	C01	74942
HV0806	382022	114	744	6262h	0	139	424592	75138	143333	202170	100	76882
HV0807	· -			45144		133	424129	75154	143487	505169	105	78922
HAURÙB	382060	114	720	6215b			424763			922202	67	79006
HV0809	382077			-		125	424795	75193	143988	202251	26	79029
HV0810	382095					123	424829	75216	144263	202277	65	79141
hv0811	382122						489883				-60	79169
Hv0812	382157			-		1 1 1	424956			505348	-17-	19263
HV0813	382186					110	42500m			202411	<b>-</b> 109	79359
HVORTA	182241						425113			202491		79646
HV0815	382240			SONIS	-		465125	_		202495	* t	79717
HVOBIN	362245								-	505497		79514
HVOR17	382580		-	95944			475146					79454
HV0818	385583			5334n			425213			295205	_	79475
HV0819	382300	114	•	n 36-4 Y						202576		79550
HVORED	362317							-		505603	1549	79688
HV0821	382334	•	77					•		202028	1030	79850
428014	382344	114	51	44624	U	146	425319	70128	143689	505005	1000	79972

FRE Ut LIST

PROFILE #9 HARETY VALLEY GEAVIEY DATA

		LUNG. ELE					ORSV GRAV	THEC	FAA	CBA +1600
•	*								~ - ~ ~ _	
HV0901		114 279 65				750391				79165
HV0902	381642	114 265 648	328 1	2334	24311	758581	41253	201013	550	74775
HVQQAZ	381659	114 250 64	Q = ()	2154	24033	755791	41612	291536	340	7870B
Hv0904	381664	114 220 63	748 O	1934	24055	759221	41812	201645	154	72609
HV0905	381673	114 190 63	48 0	1800	24072	754661	42124	01659	-107	78538
PA0309	381636	114 162 628	SOB U	1714	7.0045	760001	423178	201077	-245	76504
HV0907	381599	114 134 62	158 0	1644	24125	760461	42490	201697	-432	78433
HV0908	381720	114 115 62	138 0	1504	24162	750721	427442	757109	-599	78403
P0909	381742	114 100 612	19A 0	1694	24203	760931	43237	201759	- 84C	78425
MV0910	381764	114 Bo 619	168 0	1714	24245	701121	43401	201792	USF-	74426
HV0911	381761	11+ 54013	19 <b>T</b> 0	1004	24241	761501	43257	201788	-017	72434
HV0912	381768	1135958 629	128 0	1014	74258	762981	42093	201796	نے وہ کے 🕶	74575
HV0913	381805	1135860 645	52B 0	1754	24331	704391	21477	201-52	757	79023
HV0914	381851	11 15790 665	0 45	1004	24425	70530	41/451	25495c	2154	70000
HV0915	381889	1135726 64	148 0	2124	24494	740291	41047	201775	3497	00376
HV0916	381917	1135644 69	10 HP2	2544	24549	757471	40719	202116	3425	80544
HV0917	381983	1135572 720	N 14 (1)	34.14	23074	768481	38345)	202113	4034	F0101
HV0918	382063	1135528 700	esp v	2874	24825	769071	36460	200030	500h	79845
HV0919	385000	1135435 71	55 W	2591	द्धपुष्ठ	770901	390000	31.22.28	3706	79759
HV0920	382105	1135410 705	554 0	20×4	24908	770761	505257	20262	3034	74742
₽V0451	362112	1135382 70	\$50 0	1434	54925	771171	397426	202302	3053	79052
<b>FVU355</b>	382120	1135356 690	3 K ()	4014	<b>८</b> वच रत	771541	40154	202314	3.57K	24461
HV0923	382129	1135331 69	$\alpha = \epsilon \alpha$	1864	24956	771901	404797	202327	3191	70404
P260A4	382140	1135308 686	<i>t</i> ) +44.	1704	24978	772231	-17697	202343	3930	74795
HV0925	385125	1135284 68	0 54-6	1714	52001	772571	411287	202361	2917	79842
FV0926	382167	1135265 679	58 0	-	_	772841			2744	19012
HV0927	382183	1135246 67				773101			2074	79951
HVN9Z8	382201	1135227 hn				173371			5,247	79976
HV0920	382216	11352116688	10.1	1594	25173	773591	451419	505422	5034	79982

END OF LIST

